

# Methane on Mars II workshop

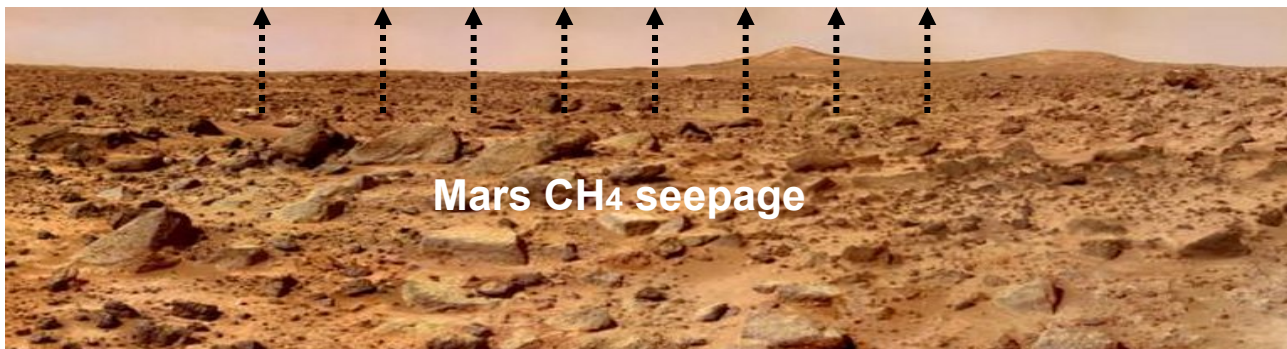
## “Gas seepage”

Giuseppe Etiope

(Seepage WG: G.Etiope, P. Beckett, J. Blank, P. Chen, R. Klusman, D. Oehler, L. Shen)

The occurrence of  $\text{CH}_4$  in the Martian atmosphere may imply active geologic sources, i.e. gas emission structures in the Martian soil and subsoil

= gas seepage, a process well known on Earth, should exist on Mars.



# The concept of “gas seepage” on Mars

How does seepage concept translate to the Mars context?

wrongly considered, erroneously associated to volcanic emission only, in:

[Icarus 178 \(2005\) 487-492](#)

A sensitive search for SO<sub>2</sub> in the martian atmosphere: Implications for seepage and origin of methane

Krasnopolsky V.A.

First systematic and geologic-geochemical discussions in:

[Planetary and Space Science 59 \(2011\) 182–195](#)

Methane emissions from Earth’s degassing: Implications for Mars

G. Etiope<sup>a,\*</sup>, D.Z. Oehler<sup>b</sup>, C.C. Allen<sup>b</sup>

[Icarus 224 \(2013\) 276–285](#)

Low temperature production and exhalation of methane from serpentinized rocks on Earth: A potential analog for methane production on Mars

Giuseppe Etiope<sup>a,b,\*</sup>, Bethany L. Ehlmann<sup>c,d</sup>, Martin Schoell<sup>e</sup>

NO  
VOLCANOES

# Outcomes of the 1<sup>st</sup> workshop

Gas seepage introduced: - basic concepts and observational data on Earth  
- potential seepage on Mars

Gas seepage on Mars:

- can be evidenced by specific **surface manifestations (macro-seepage structures)** over faults and fractured rocks, as observed on Earth (*circular depressions, polygonal fractures, mounds, mud volcanoes*).
- can be in the form of invisible diffuse exhalation from the ground (**microseepage**)
- can be detected only through specific procedures/methods: measurements in the atmosphere, a few cm above the ground (as performed by Curiosity) may not be effective in revealing any seepage (trivial CH<sub>4</sub> recorded by Curiosity cannot be accepted as evidence of lack of subsoil processes generating methane)
- surface gas geochemical techniques, similar to those that allowed the discovery of seepage and subsoil hydrocarbon reservoirs on Earth, must be considered (soil-gas, accumulation chambers, surface mineralogical alterations).

## Take home message

Geologic CH<sub>4</sub> on Mars should be searched, preferably above or near faults or at apparent mud volcanoes, in the regions with olivine bearing or sedimentary rocks, ideally by drilling into the soil, or using accumulation chambers on the ground

# Objectives of the 2<sup>nd</sup> workshop

## More detailed discussions on.....

### **1. Macro and microseepage on Earth**

typical soil-gas concentrations and flux values, detection methods, indirect methods (soil-gas, closed-chamber, drilling, instrumental requirements. What can be applied on Mars?)

### **2. Potential seepage structures and manifestations on Mars**

recognition by high-resolution images (e.g. HiRISE, CaSSIS, land-based cameras)

### **3. Seepage proxies:** carbonate cement, secondary alterations of minerals.....

### **4. Meaning of methane/ethane ratio in seeping gas**

review of methane genetic mechanisms; post-genetic alterations during seepage, meaning (and ambiguity) of C<sub>1</sub>/C<sub>2</sub> ratio (expected to be measured by ExoMars 2016)

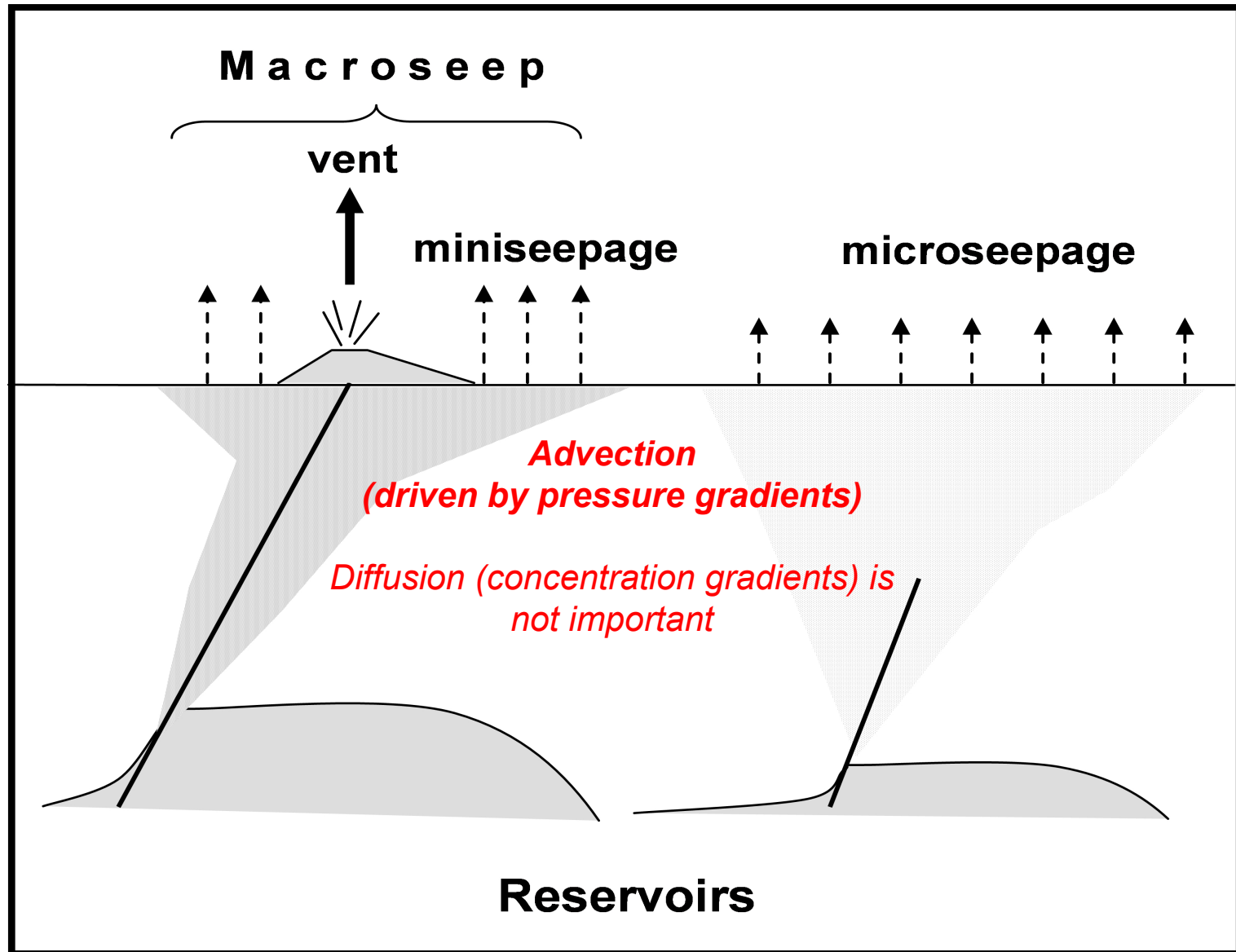
## Proposed deliverable

- Report in a format suitable for submission to a peer-reviewed journal
- Integration with the overall results of the workshop, for a comprehensive paper (e.g., EOS)

**We can propose the search for gas seepage as one of the guiding scientific goals for Mars exploration in the 2020's.**

# What is gas seepage

visible or invisible, focused or diffuse over large areas



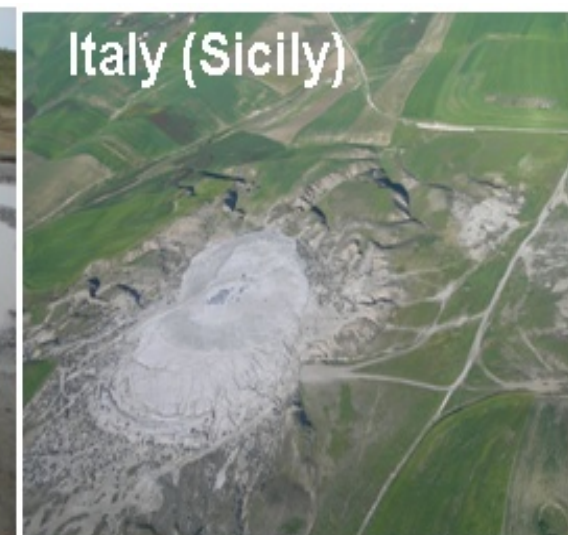
# Gas seeps and “eternal” fires

Release from 1 to 1000 ton CH<sub>4</sub> per year



# Mud volcanoes

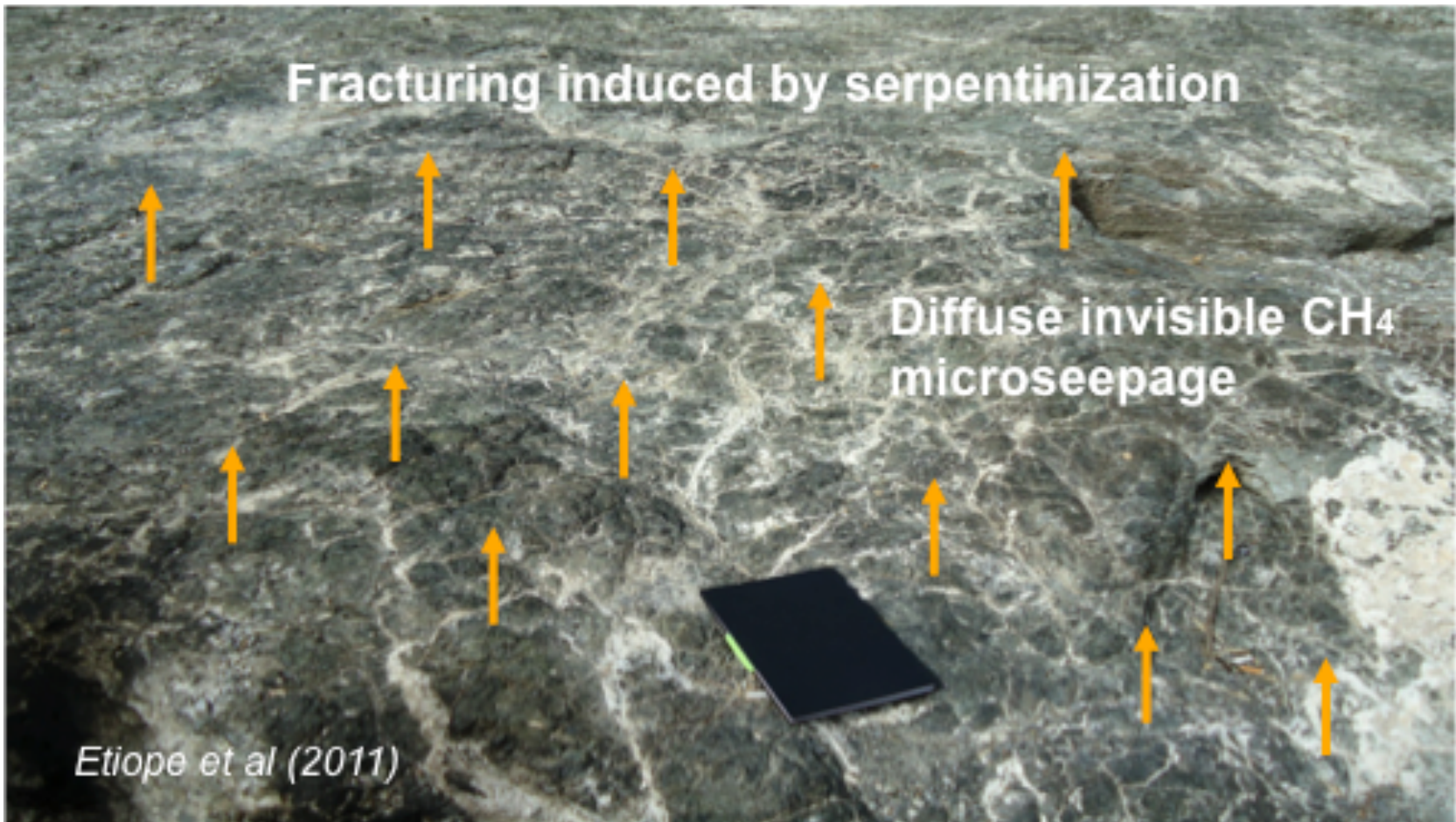
Release from 1 to 500 ton CH<sub>4</sub> per year



**Sedimentary volcanism, 3-phase system : gas-water-sediment**

# MICROSEEPAGE IN OLIVINE-RICH ROCKS (PERIDOTITES)

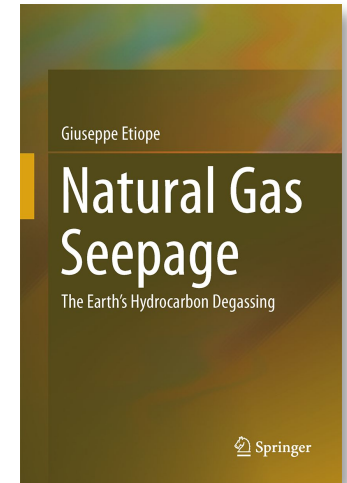
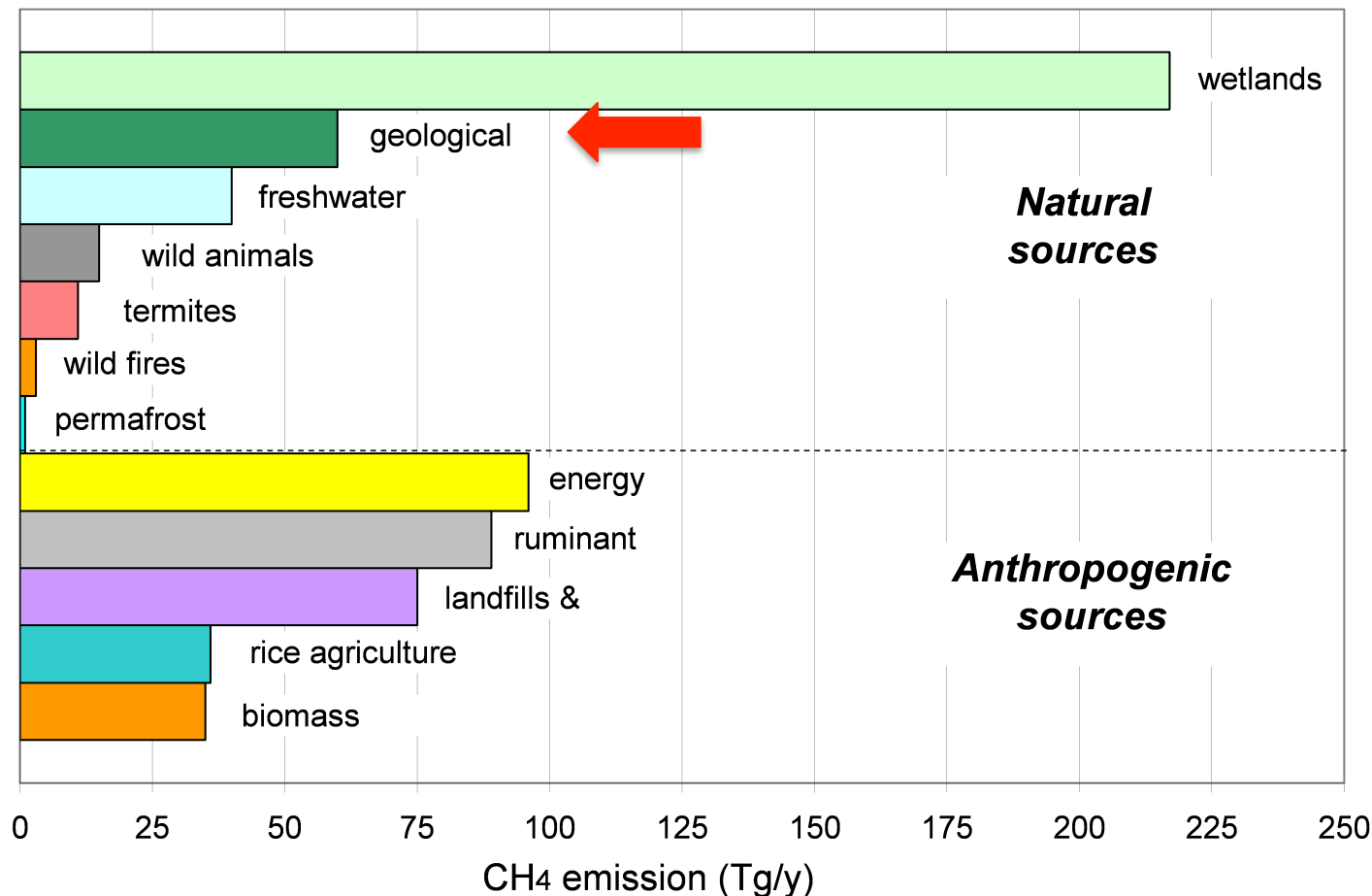
from 1 to  $10^3 \text{ mg m}^{-2} \text{ day}^{-1}$



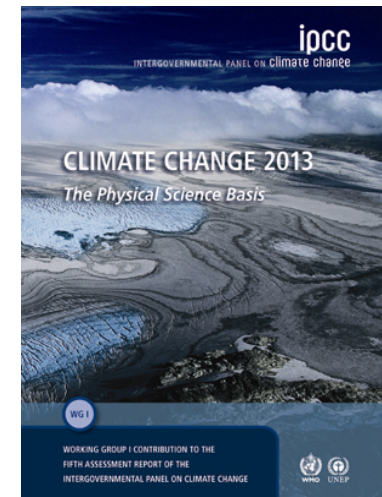
# Geological emissions in the global CH<sub>4</sub> budget

**2<sup>nd</sup> natural CH<sub>4</sub> source**  
**10% of total CH<sub>4</sub> source**

*Etiope and Ciccioli 2009 (Science)*  
*Etiope, 2012 (Nature Geosci.)*



*(Etiope, 2015)*

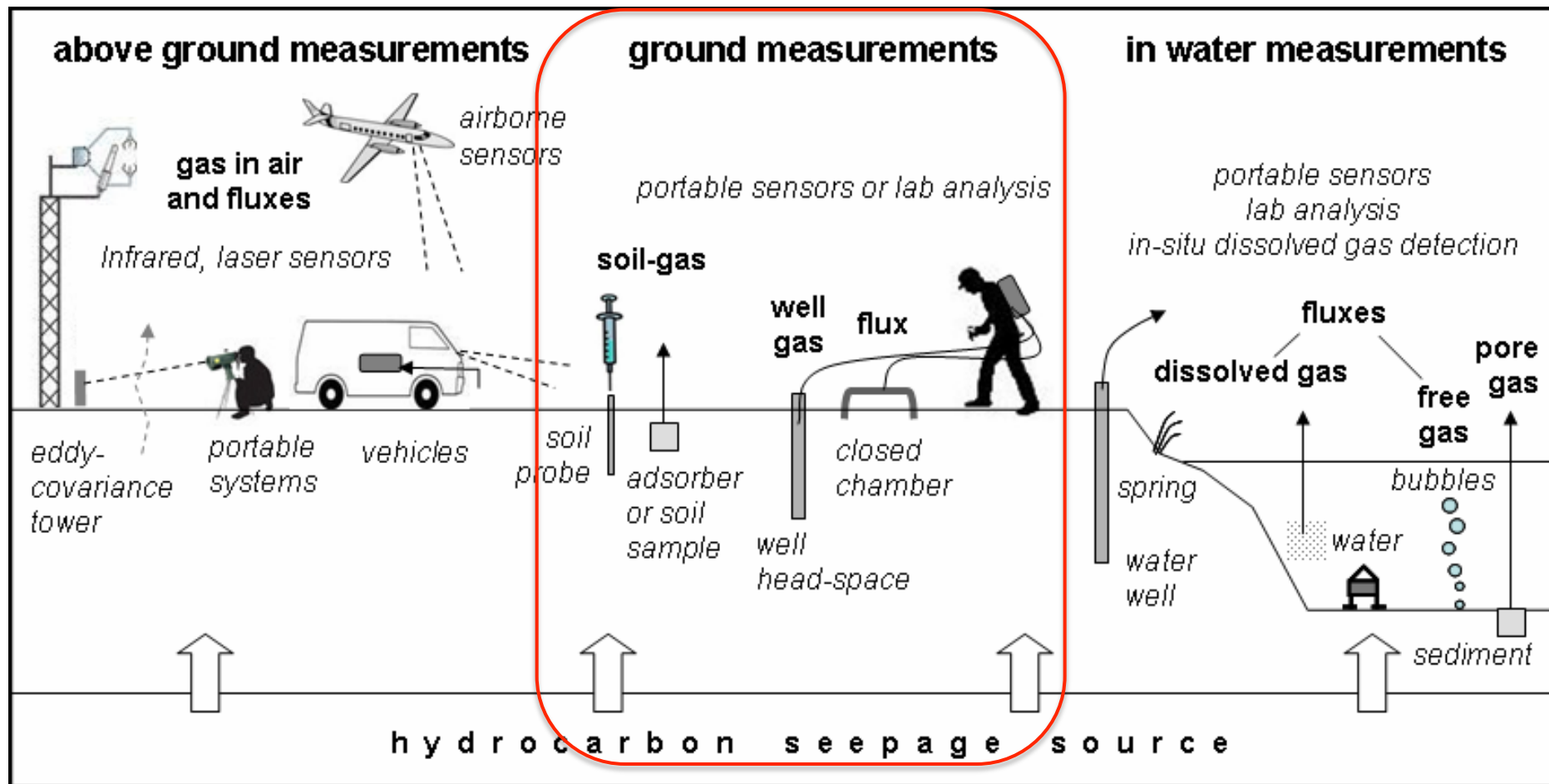


*(IPCC, 2013)*

# How to detect and measure gas seepage

Effective only for  
significant seepage

Effective also for  
very low seepage



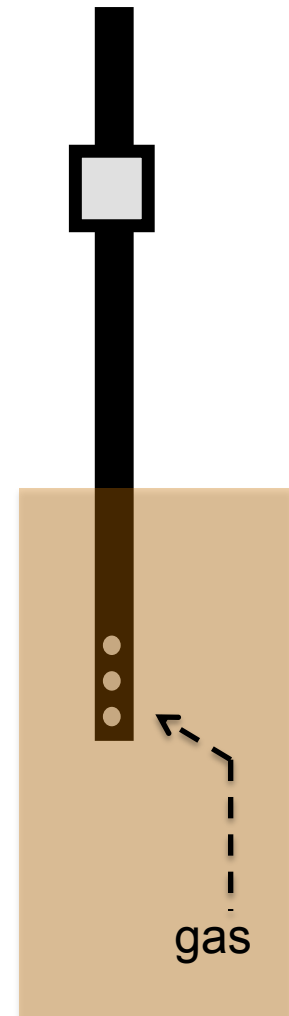
approach to be  
adopted for Mars

*Etiope (2015)*

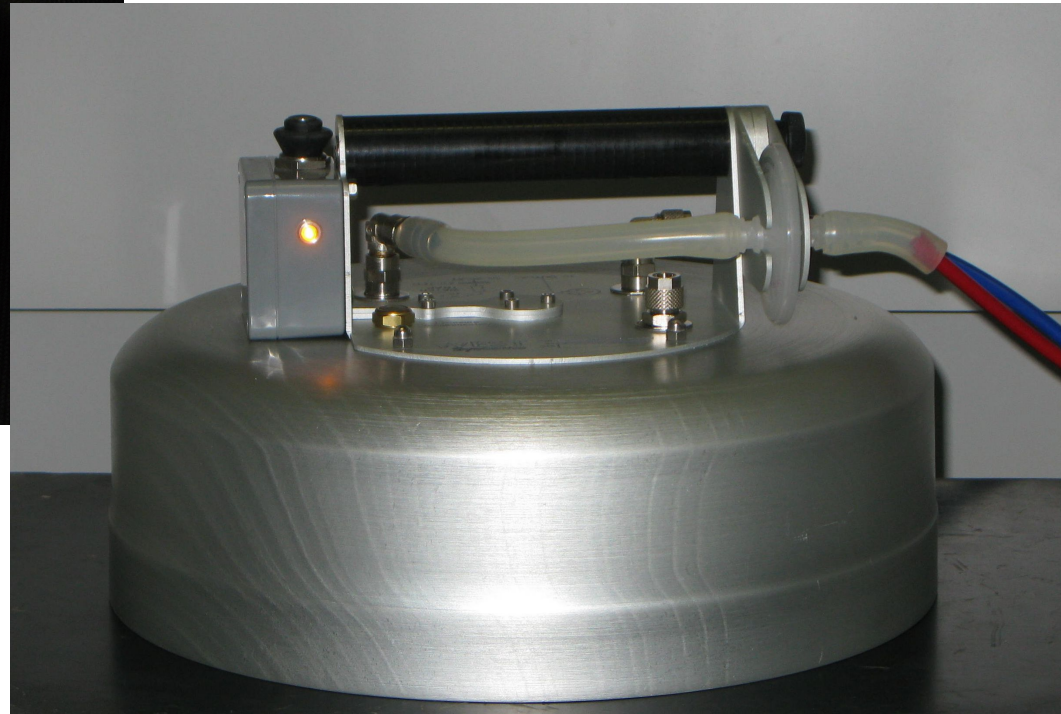
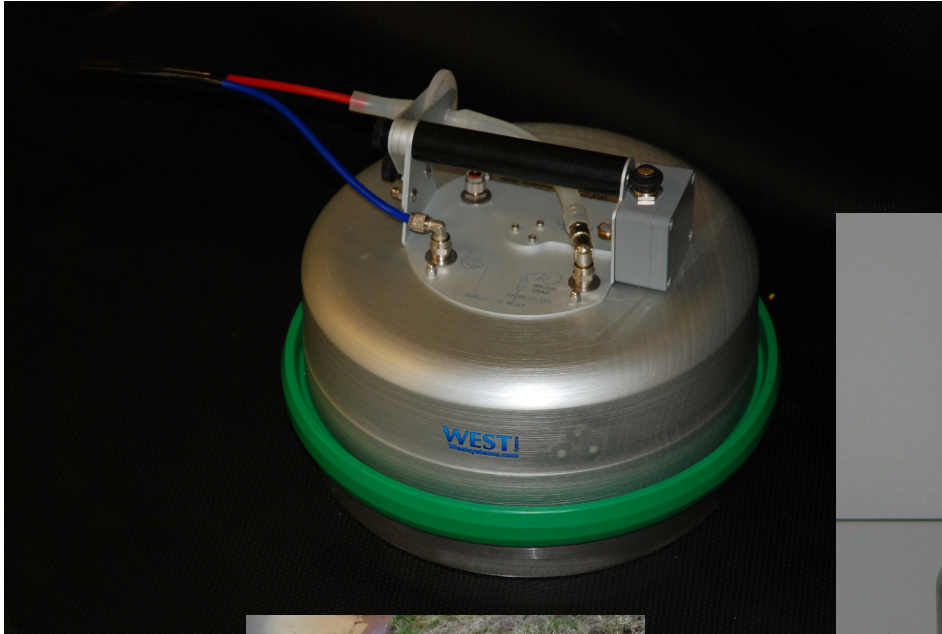
# Shallow drilling into the soil

## *soil-gas probes*

Sampling gas in soil-air at  
depths of 50-100 cm



# CLOSED-CHAMBER SYSTEM for microseepage



# Widely used for soil-respiration, gas fluxes from wetlands, rice paddies and permafrost.

Gas flux  $Q$  is expressed in terms of  $\text{mg m}^{-2} \text{ day}^{-1}$  by the eq.:

$$Q = \frac{V_{\text{FC}}}{A_{\text{FC}}} \cdot \frac{c_2 - c_1}{t_2 - t_1} \quad \left[ \frac{\text{mg}}{\text{m}^2 \cdot \text{d}} \right]$$

$V_{\text{FC}}$  ( $\text{m}^3$ ) chamber volume

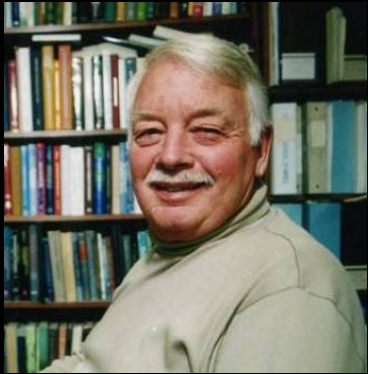
$A_{\text{FC}}$  ( $\text{m}^2$ ) chamber area

$c_1 - c_2$  ( $\text{mg}/\text{m}^3$ ) methane concentrations at times  $t_1 - t_2$  (days).



Photo: Charlotte Sigsgaard

# First applications in geology



**Prof. Ronald  
Klusman**

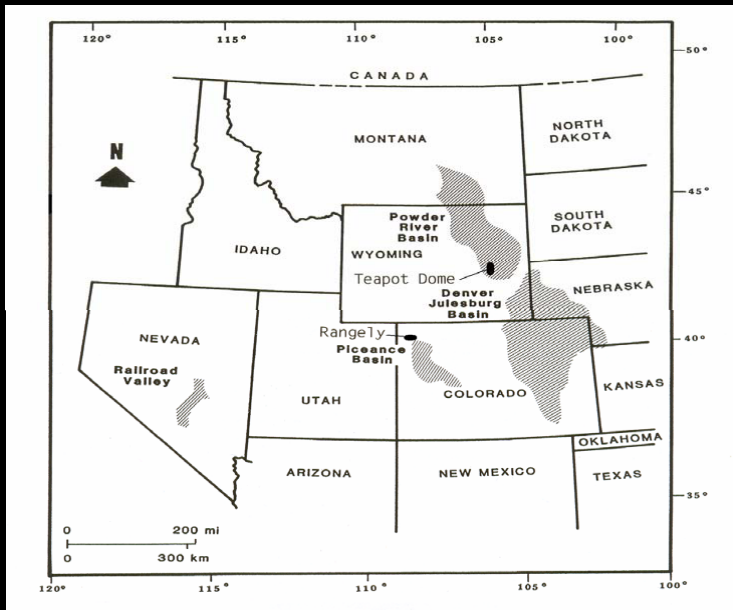
**US Colorado  
School of Mines**

Klusman et al 2000. J. Geoph. Res. 105D, 24,661-24,670

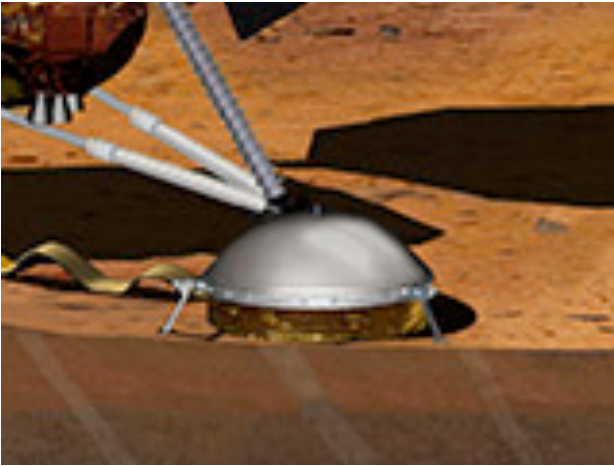
Klusman et al 2000. Geothermics 29, 637-670

Klusman 2003. Applied Geochem., 18, 1825-1838.

Klusman 2006. Applied Geochem., 21, 1498-1521.

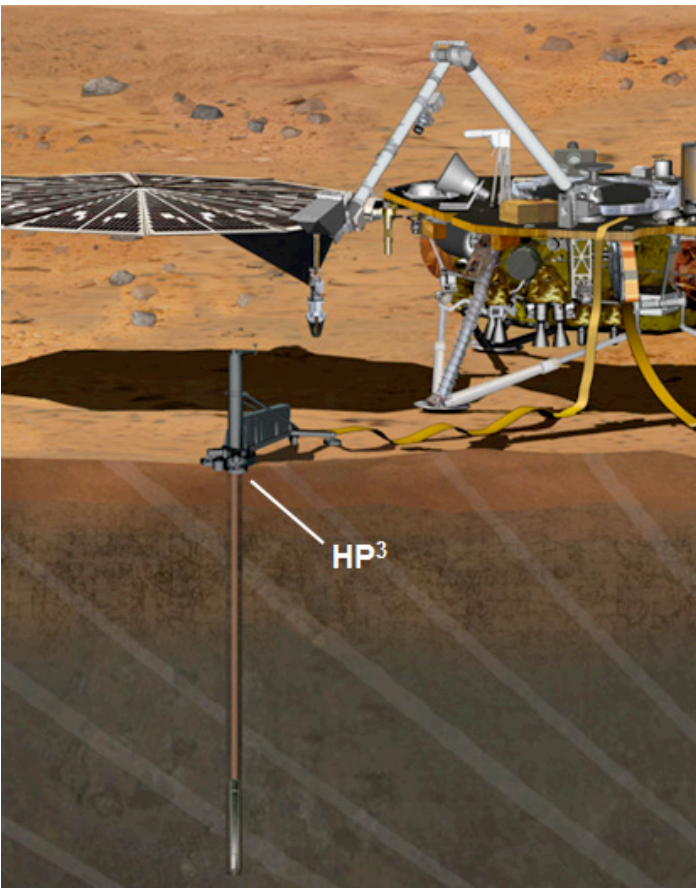


***INSIGHT mission*** (2018, NASA)  
*landing site: Elysium Planitia*



**SEIS** seismometer

A similar arm could be used for positioning a closed-chamber



**HP3** (Heat Flow and Physical Properties Probe)

CONNECTING A GAS SENSOR TO THESE  
PROBES WOULD BE A GREAT  
OPPORTUNITY TO RELIABLY DETECT  
METHANE SEEPAGE

**Can KISS support the development  
of such a concept and technology?**

*(prototype design, development, involvement  
of robotics companies)*

# POTENTIAL SEEPAGE ON MARS

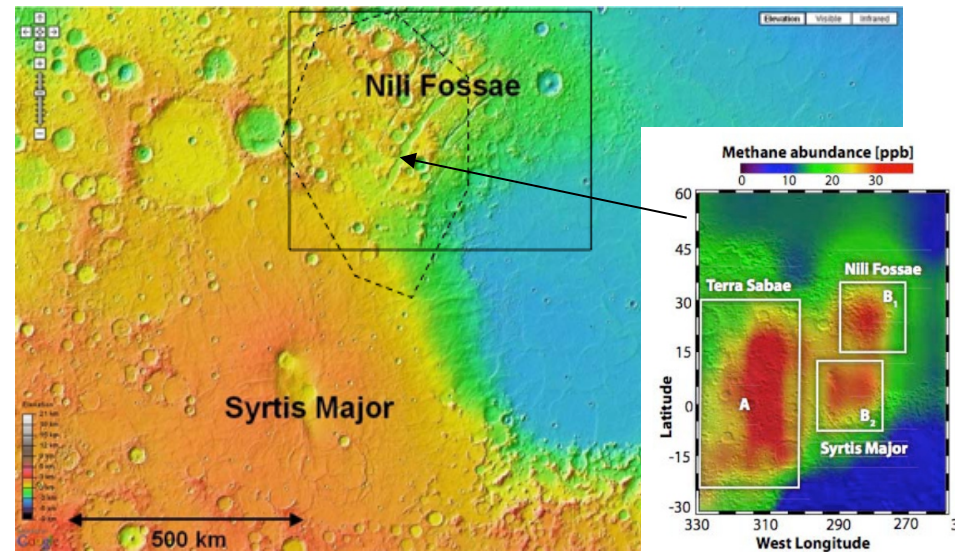
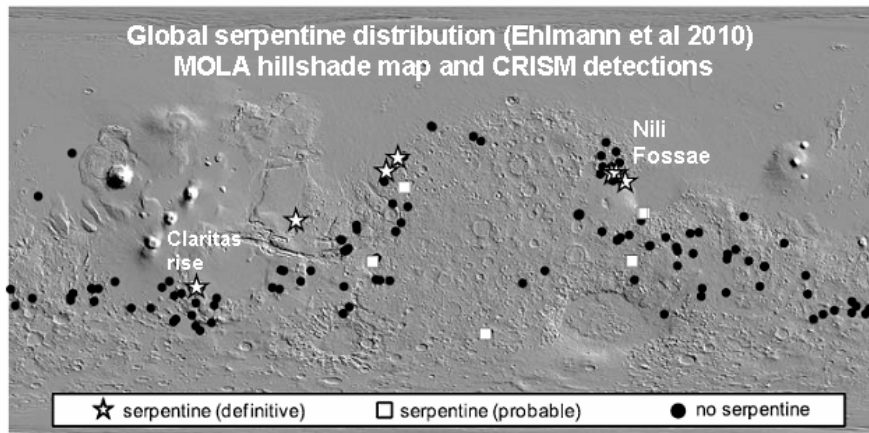
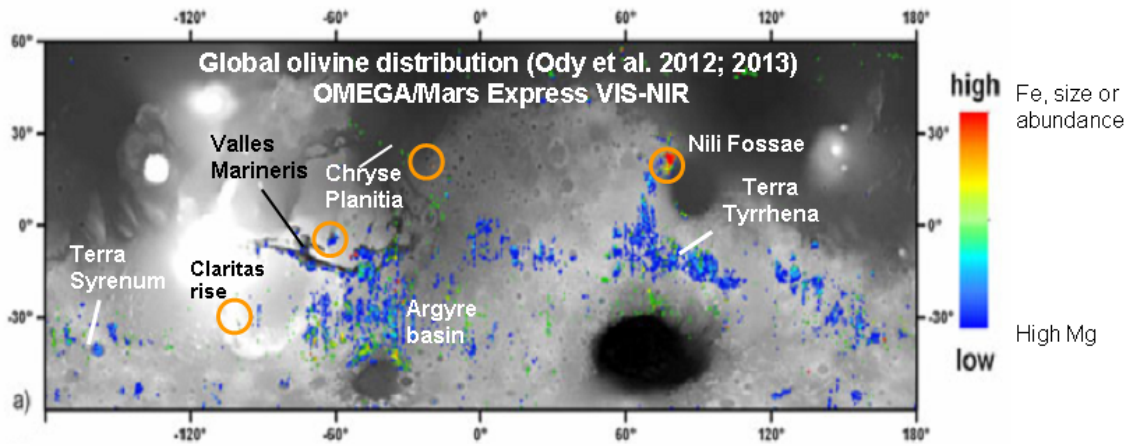
where on Mars are the best chances of finding methane?

## Analog seepage sites

faulted/fractured ultramafic/serpentinized rocks  
faulted/fractured sedimentary basins (mud volcanoes, mounds)

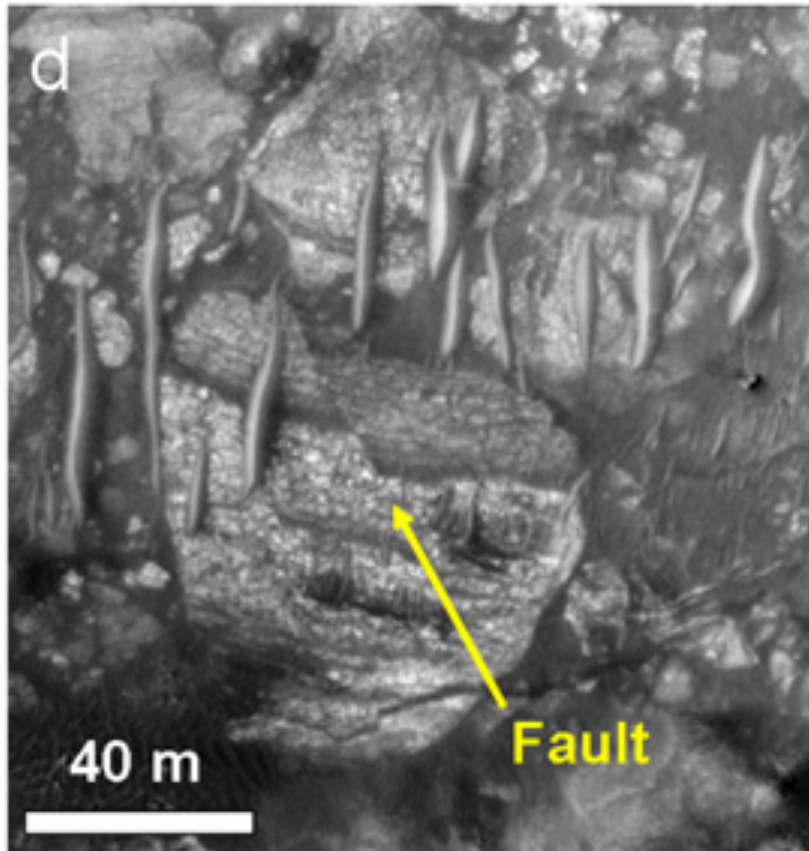
# Olivine-rich and serpentinized areas on Mars

Serpentine occurs in Mars' ancient Noachian terrains, Nili Fossae, Syrtis Major, Claritas Rise

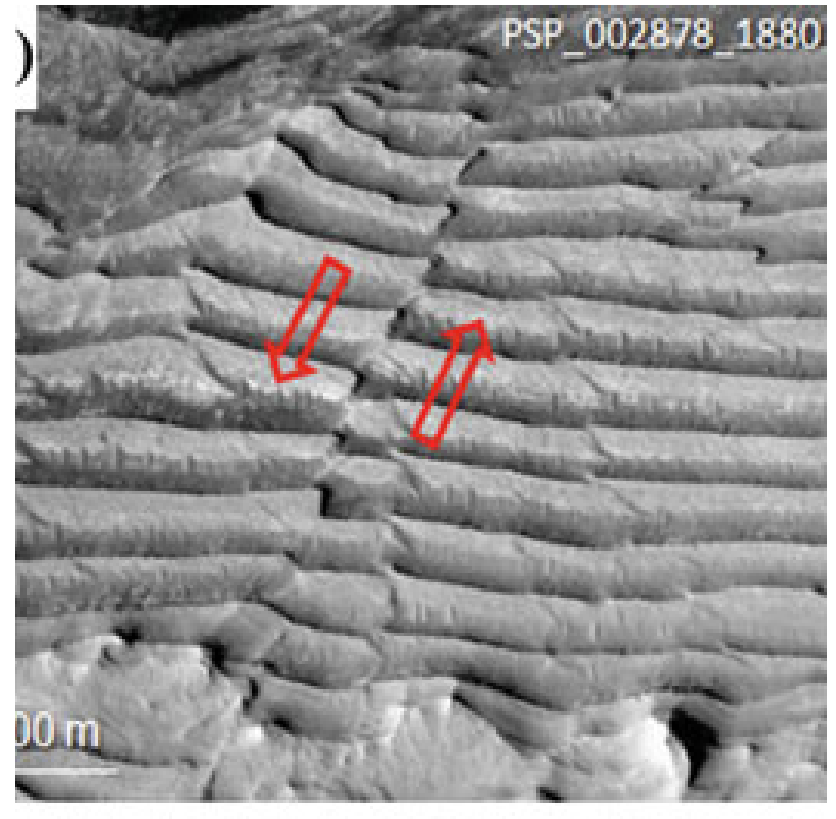


30,000 km<sup>2</sup> olivine-rich outcrop (*Hoefen et al 2003*)

# FAULTS



Fault at Nili Fossae, from  
PSP\_006923\_1995 (19.381N, 76.421E)  
*Wray and Ehlmann (2011)*



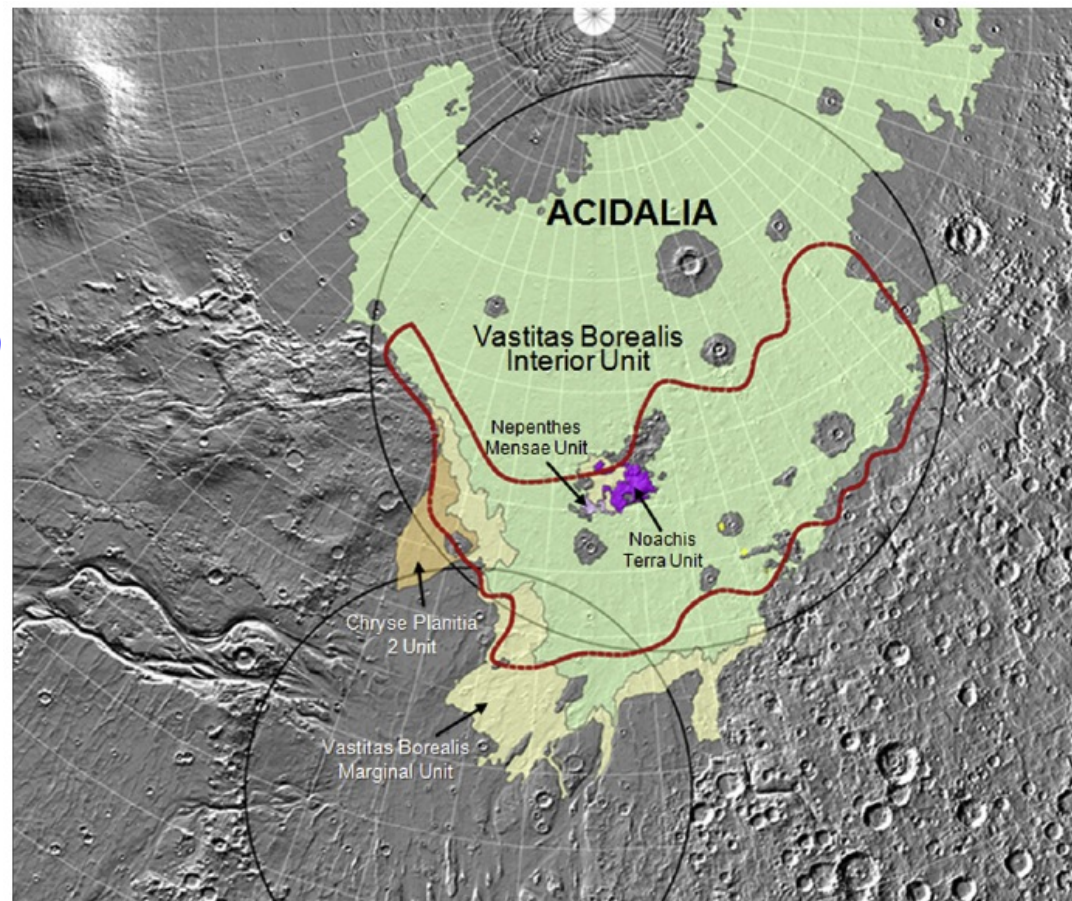
Arabia Terra  
(*Etioppe et al. 2011*)

## Potential mud volcano-like seeps

Candidate mud volcanoes reported from Utopia, Isidis, northern Borealis, Scandia, Chryse–Acidalia region ([Davis and Tanaka, 1995](#); [Tanaka, 1997, 2005](#); [Tanaka et al., 2000, 2003, 2008](#); [Farrand et al., 2005](#); [Kite et al., 2007](#); [Rodríguez et al., 2007](#); [Skinner and Tanaka, 2007](#); [Allen et al., 2009](#); [Oehler and Allen, 2009](#); [Skinner and Mazzini, 2009](#); [McGowan, 2009](#); [McGowan and McGill, 2010](#))

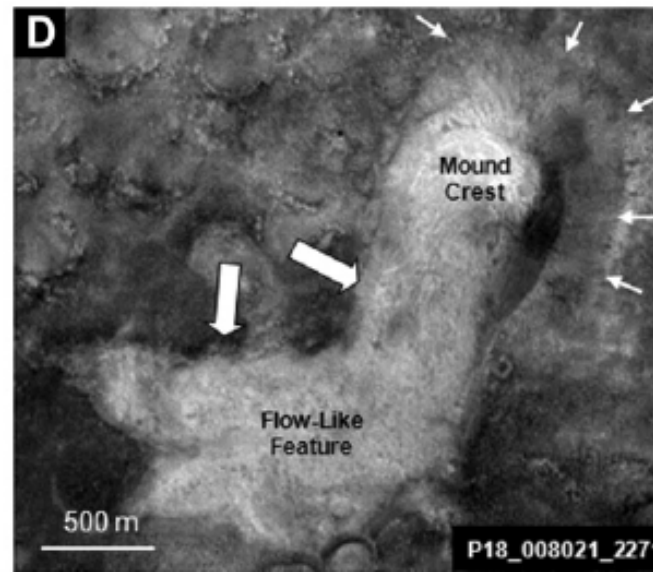
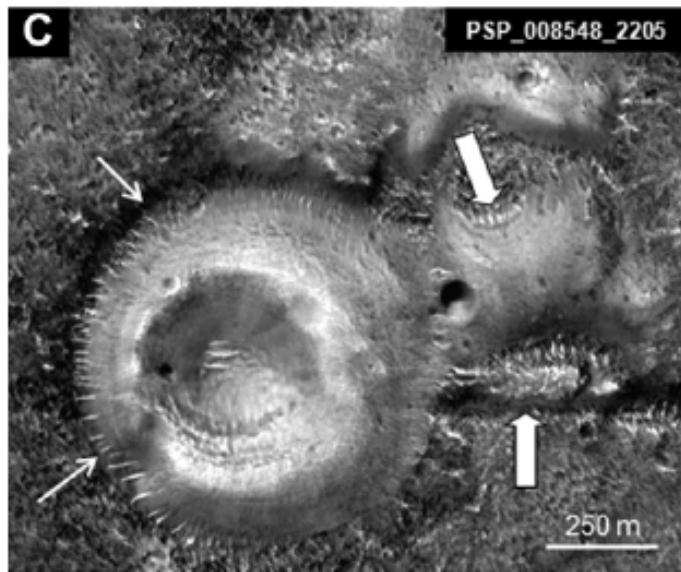
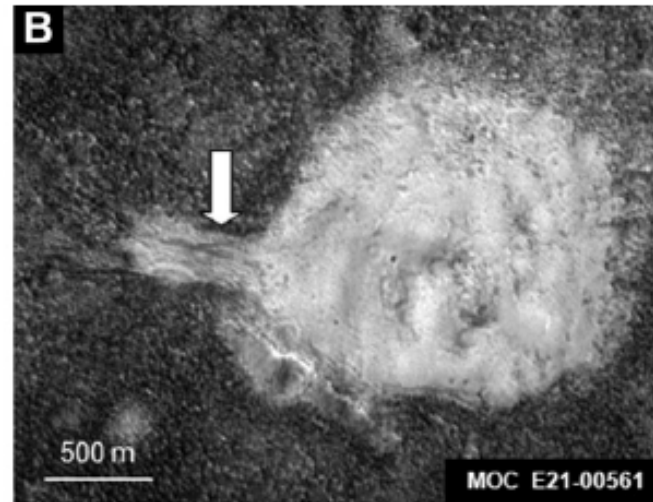
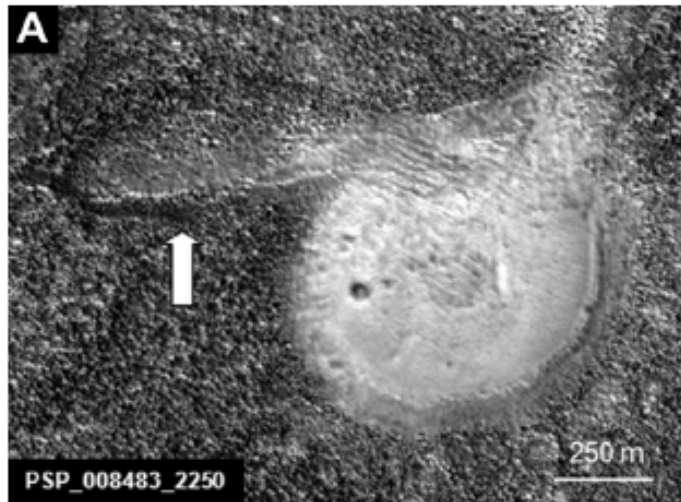
>40000 estimated

(18000 mapped)  
in Acidalia Planitia  
([Oehler and Allen, 2010](#))



# Potential mud volcano-like seeps Acidalia Planitia

(*Oehler and Allen, 2010; Etiope, Oehler, Allen 2011*)



# INTERPRETING MOLECULAR COMPOSITION OF GAS

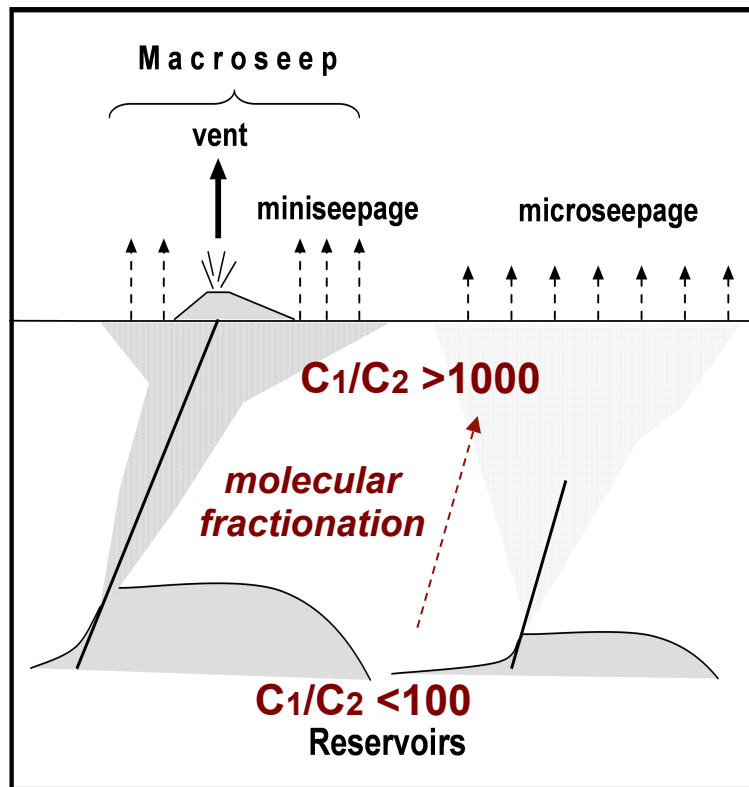
## THE MEANING OF METHANE/ETHANE RATIOS

(expected to be measured in the ExoMars 2016 mission)

is it a reliable indicator of gas origin?

It is generally assumed that  $C_1/C_2$  is a good indicator of methane origin, microbial ( $C_1/C_2 > 1000$ ) or thermogenic or abiotic ( $C_1/C_2 < 1000$ )

But this is true only if gas is sampled/detected at the point of its origin  
*(on Mars we may only detect seeping gas)*

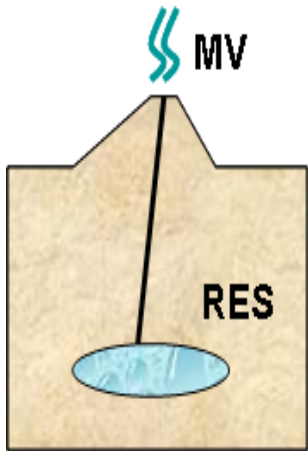


During gas migration, due to molecular adsorption on solid grains of mud/sediments seeping gas becomes dryer (more  $CH_4$ , less  $C_2+$ ) than reservoir gas.

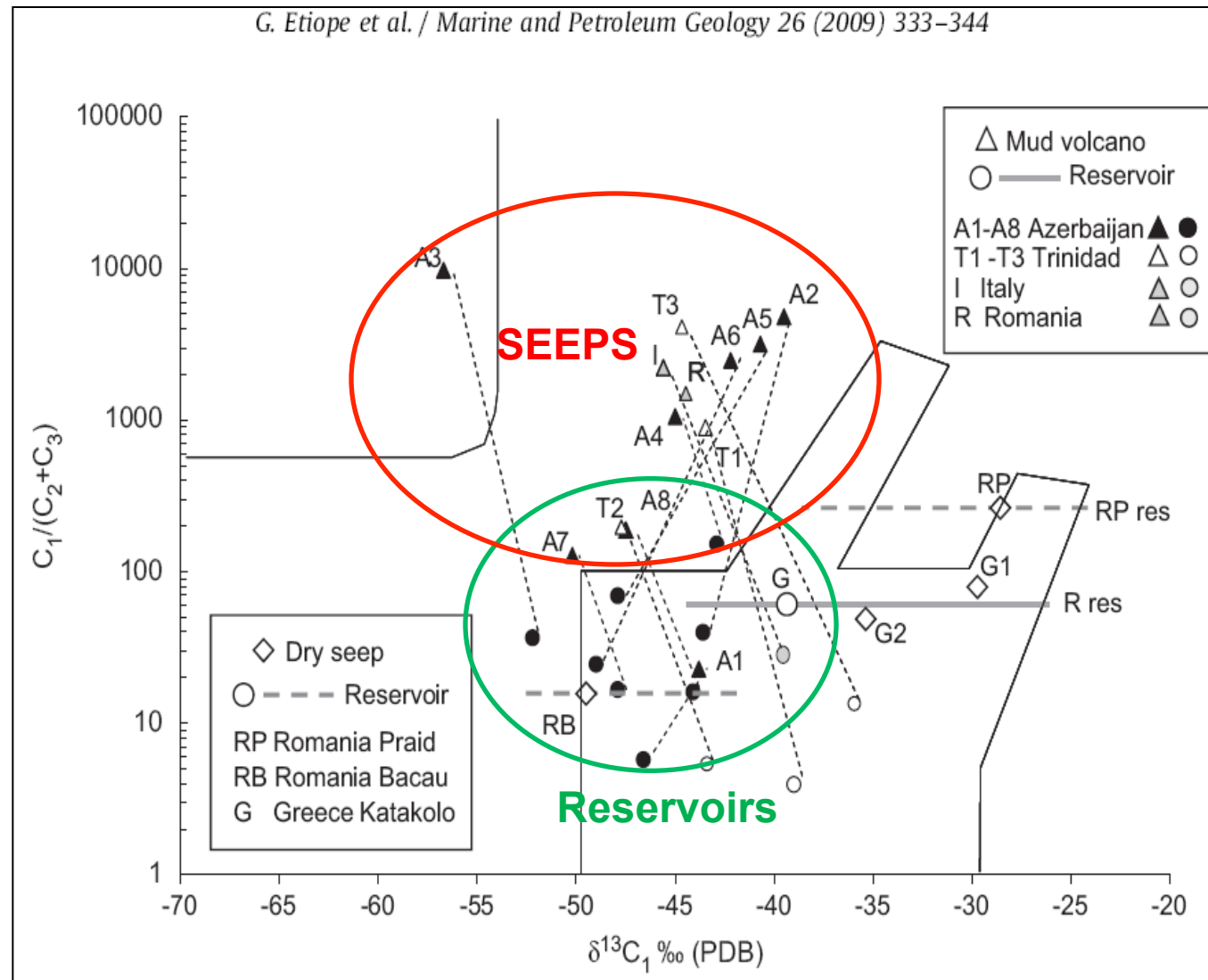
Molecular fractionation is inversely proportional to the flux or velocity of gas

# Gas migration from reservoir to surface seeps:

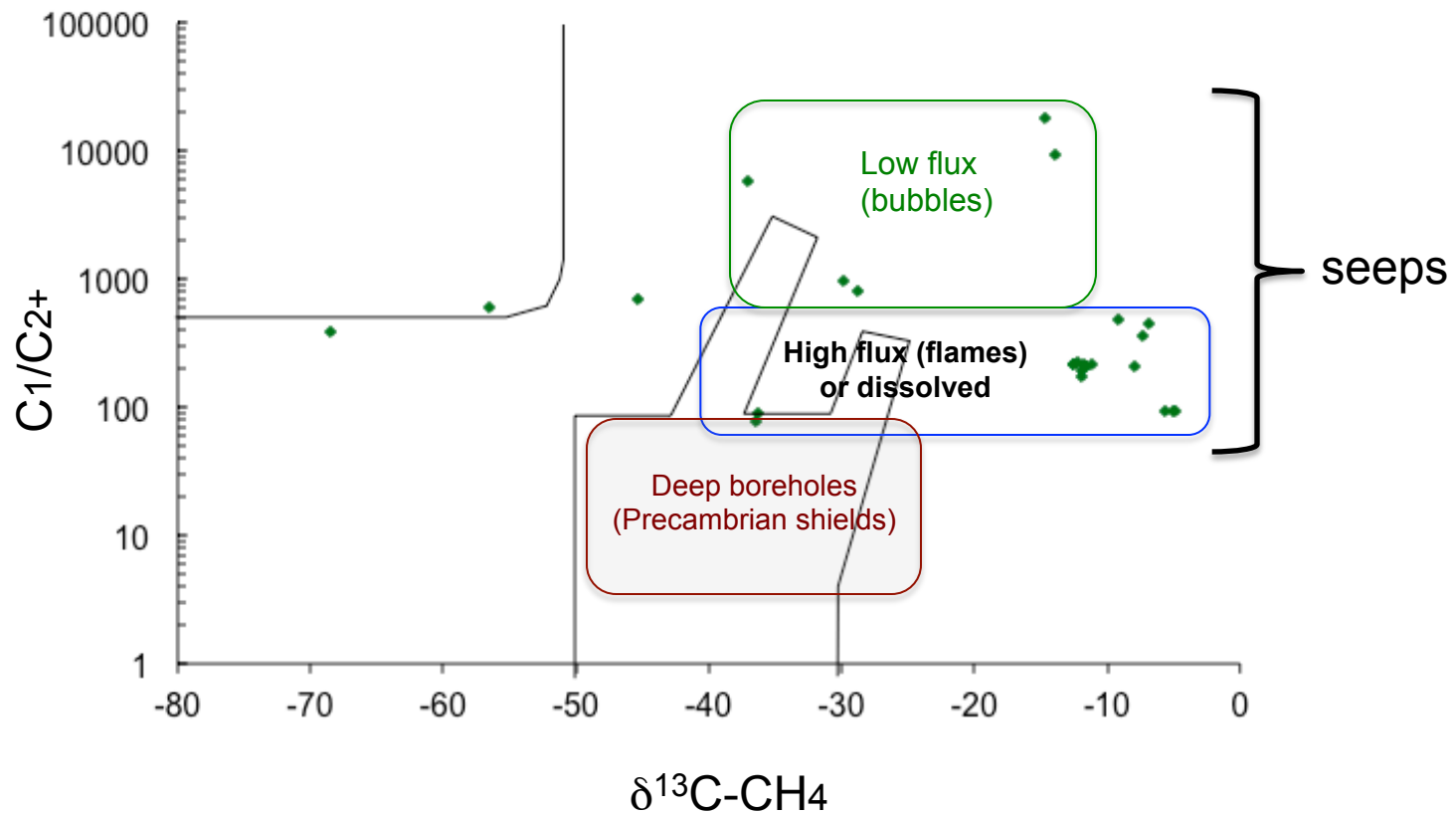
## Loss of C<sub>2</sub>+ hydrocarbons due to molecular fractionation



Analysis of  
seeps over  
corresponding  
reservoirs

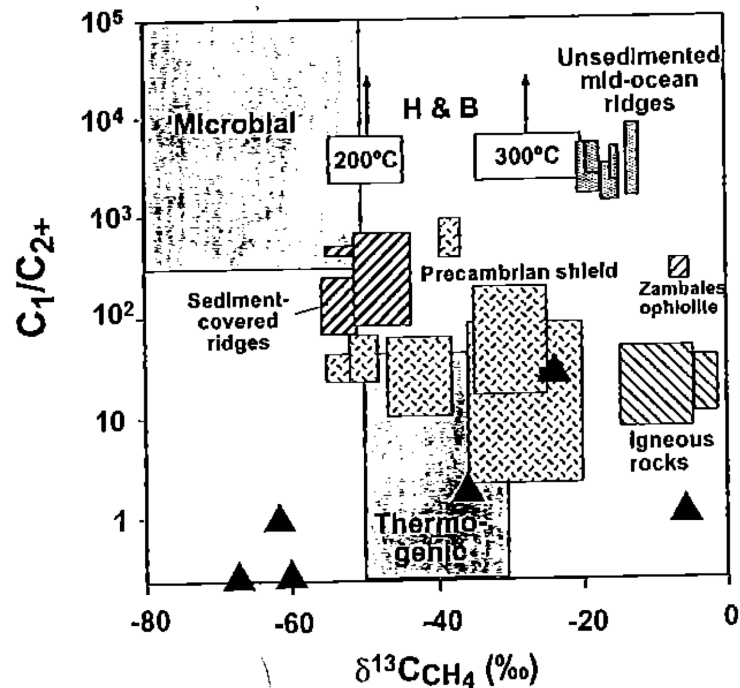


ABIOTIC GAS  
seeps (continental serpentinization sites)  
vs  
deep boreholes (Precambrian shields)



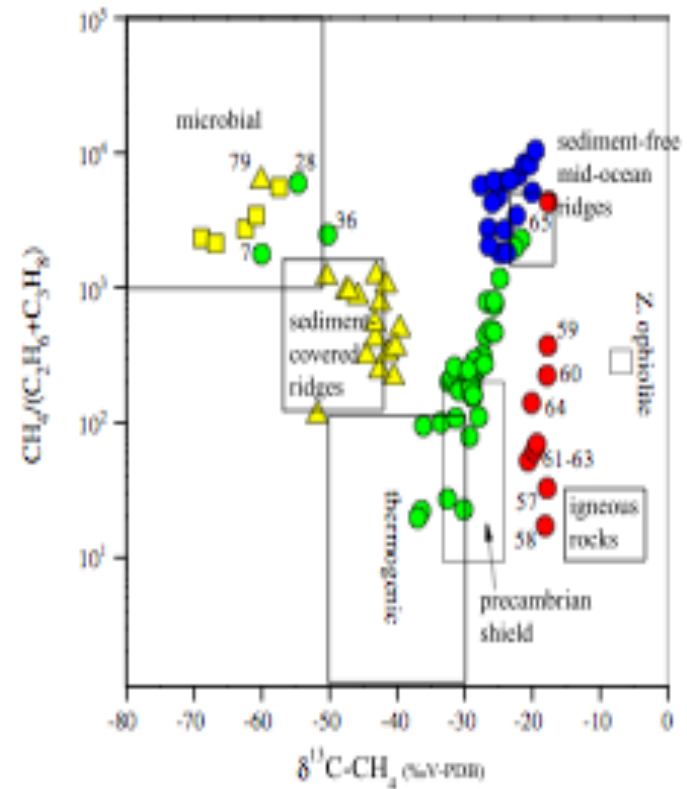
# OTHER EXAMPLES OF ABIOTIC and GEOTHERMAL GAS

McCollom & Seewald (2007)



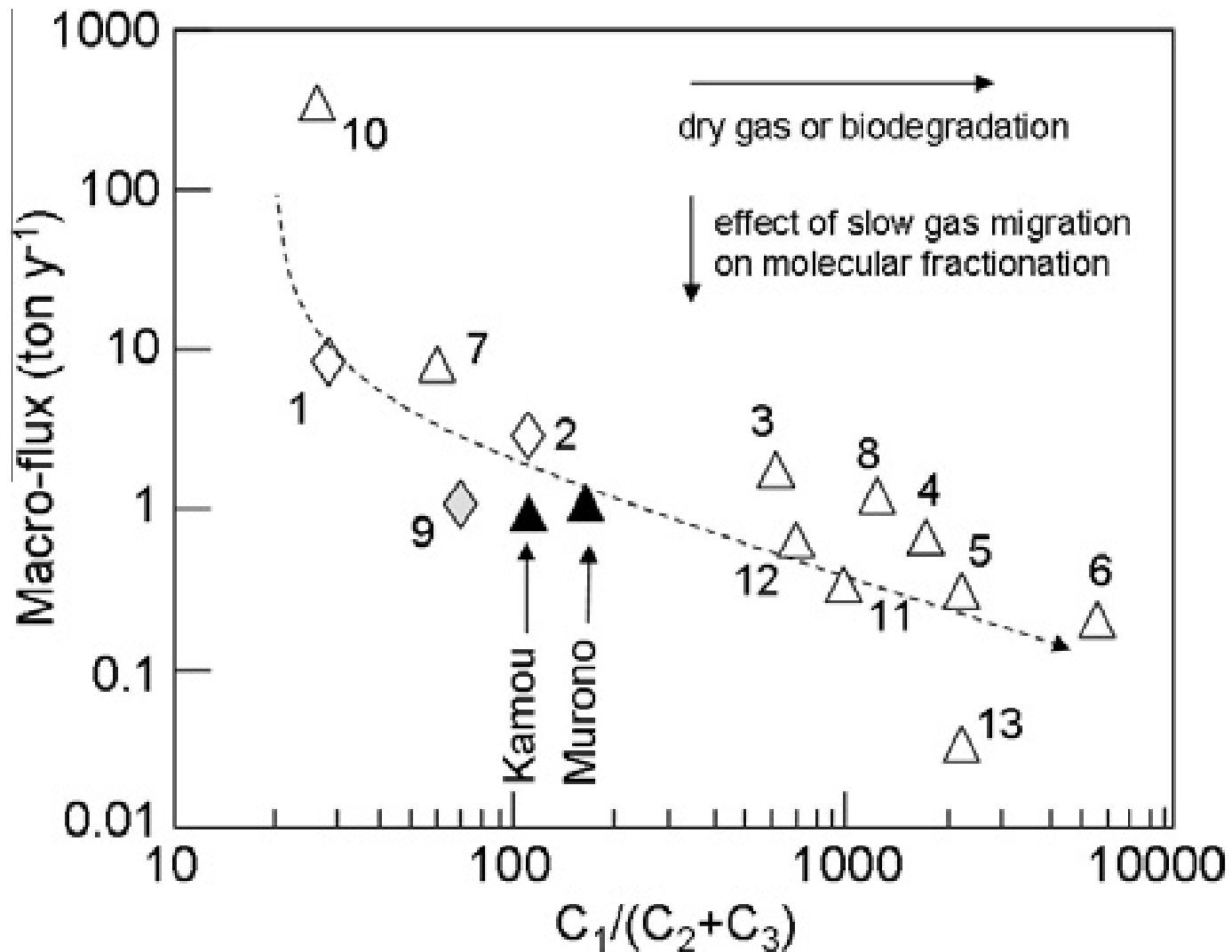
**Figure 19.** Carbon isotopic composition of methane and  $C_1/C_{2+}$  ratios for volatile hydrocarbons in high-temperature fluids from mid-ocean ridge hydrothermal systems. Shown in gray is the range of values observed for typical microbial and thermogenic hydrocarbon gases. Also shown for comparison are values for gases proposed to have an abiotic origin in Precambrian shield rocks, igneous rocks, and gases venting from a continental serpentinite (Zambales ophiolite). The fields marked "H & B" are abiotic methane generated in hydrothermal experiments of Horita and Berndt<sup>43</sup> ( $C_1/C_{2+}$  values are minimums because  $C_{2+}$  compounds were below detection limits). ▲ represent hydrocarbons formed in Fischer-Tropsch synthesis experiments.<sup>27,52,54</sup> Adapted with permission from ref 43. Copyright 1999 AAAS (<http://www.aaas.org>).

Tassi et al (2012)



**Fig. 3.**  $CH_4/(C_2H_6 + C_3H_8)$  vs.  $\delta^{13}C-CH_4$  diagram for the Italian gas discharges. Values for gases of biogenic origin (microbial and thermogenic) and from unsedimented mid-oceanic ridges, sediment-covered ridges and igneous rocks are reported (McCollom and Seewald, 2007, and references therein) for comparison. Green circle: gases from the Tyrrhenian domain; blue circle: gases from the geothermal systems of Mt. Amiata, Larderello, Latera, Manzianna and Ischia; red circle: gases from Solfatara, Vesuvio, Panarea and Pantelleria volcanic-hydrothermal systems; yellow triangle: gas discharges in the Apenninic domain; yellow square: gases from the Adriatic domain.

## Molecular fractionation inversely proportional to gas flux



*Etiope et al (2011)*

Therefore.....

**1.** If both **methane** and **ethane** will be detected, then it is likely that gas is abiotic, but we shall assume that

(a) there are not ethanogens on Mars (ethane-producing microbes exist on Earth)

(b) there is no ancient organic matter in deep sedimentary rocks that could be degraded by temperature (i.e. the possibility for thermogenic gas shall be excluded, and we shall explain why)

**2.** if **ethane** will **not** be detected, then we may hypothesize

(a) microbial gas

(b) abiotic gas molecularly fractionated

(c) abiotic gas generated at very low T (no enough energy for polymerization of  $\text{CH}_4$  molecules to form  $\text{C}_{2+}$ )

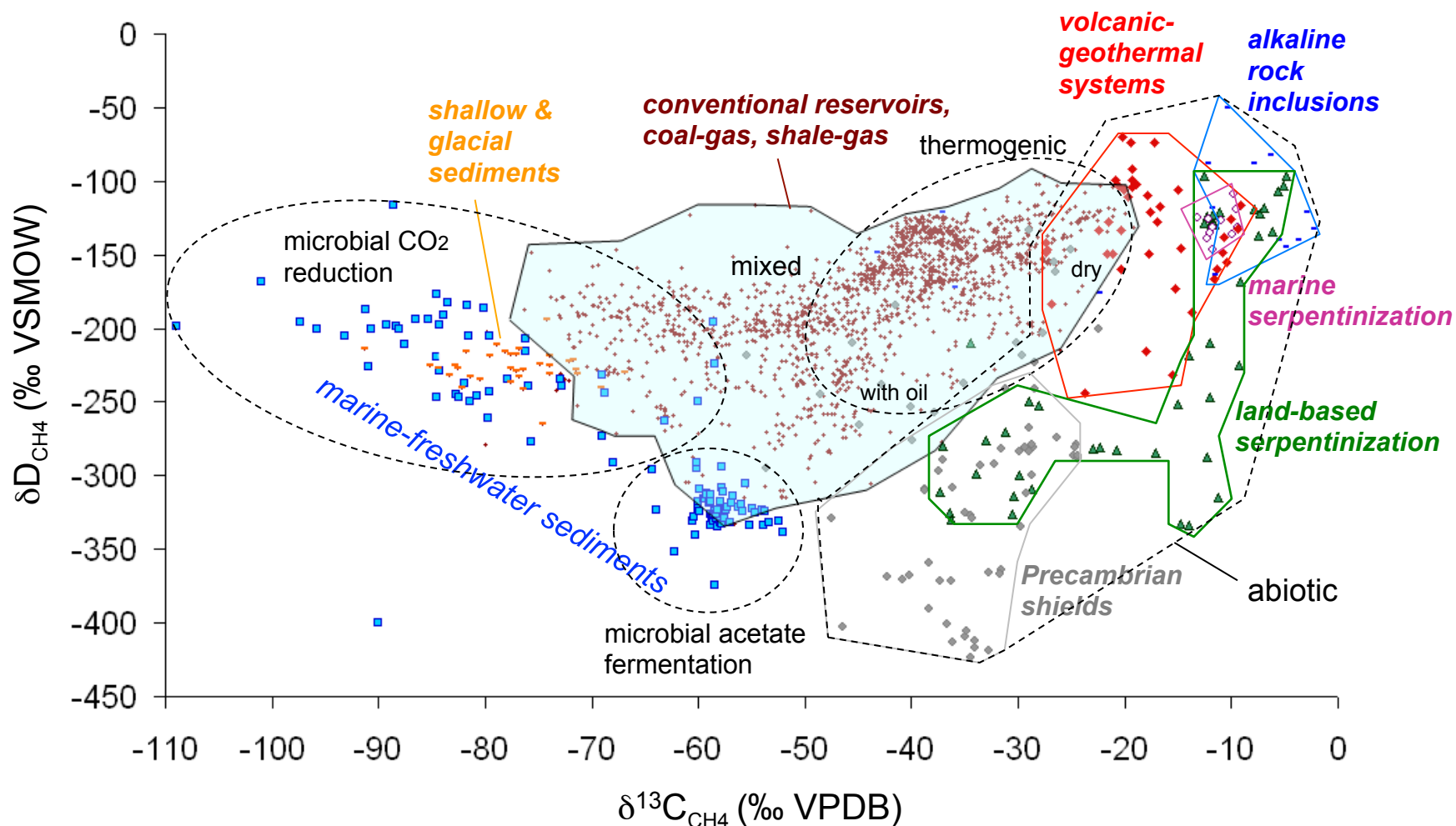
In any case, we will have a certain degree of uncertainty on the origin.

However, detecting ethane would make the interpretation a bit easier (probable abiotic gas); considering the geological framework and the features of the sampled site (mud volcano, sedimentary basin, basalt and serpentines..etc..) could help.

# INTERPRETING ISOTOPIC COMPOSITION OF GAS

The meaning of  $\delta^{13}\text{C-CH}_4$  and  $\delta^2\text{H-CH}_4$

## Updated (2015) CH<sub>4</sub> isotopic-genetic plot



*G. Etiope (2016), Encyclopedia of Geochemistry, Springer*

# Potential C–H isotopic signatures of CH<sub>4</sub> on Mars

- Martian C feedstock:**
- atmospheric fractionated CO<sub>2</sub> ( $\delta^{13}\text{C}$ : +46 ‰; [Webster et al 2013](#))
  - atmospheric unfractionated CO<sub>2</sub> ( $\delta^{13}\text{C}$  -20‰ to 0‰; [Niles et al., 2010](#))
  - magmatic CO<sub>2</sub> (Zagami meteorites,  $\delta^{13}\text{C}$ : -10 to -20‰)

$\delta^{13}\text{C}$ -CH<sub>4</sub> can be similar to that observed on Earth only if it derives from unfractionated or magmatic CO<sub>2</sub>

- Martian H feedstock:**
- atmospheric H<sub>2</sub>
  - H in minerals (meteorites)
  - subsurface waters ???
  - magma: low  $\delta^2\text{H}$ ; initial  $\delta^2\text{H}$  similar to Earth; [Boctor et al., 2003](#); [Lunine et al., 2003](#)
  - igneous rocks: olivine,  $\delta^2\text{H}$ : -60 to -280 ‰ [Gillet et al. 2002](#)
- extrem. enriched in deuterium  $\delta^2\text{H}$  up to +4000‰  
[Leshin, 2000](#); [Sugiura and Hoshino, 2000](#)  
due to atmospheric escape fractionation processes

A wide range of  $\delta^2\text{H}$  could be measured for martian CH<sub>4</sub>, far outside terrestrial variations

Martian  $\delta^2\text{H}$ -CH<sub>4</sub> values could be within the terrestrial range if the precursor hydrogen derives from primordial, unfractionated, magmatic gas or is similar to that of martian olivine.

# Conceptual summary for seepage on Mars

