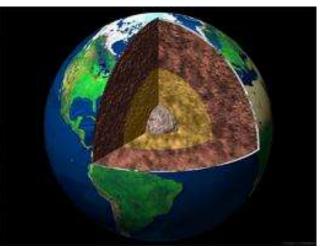
Planetary Materials and their Response to Tidal Deformation

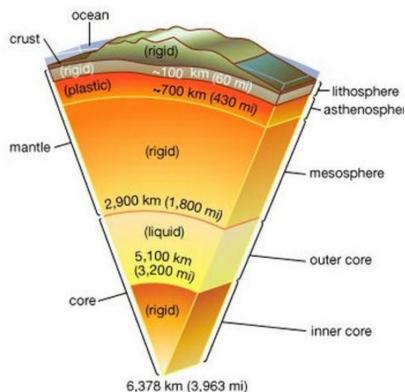
Christine McCarthy
Assistant Research Professor

Lamont-Doherty Earth Observatory
COLUMBIA UNIVERSITY | EARTH INSTITUTE

Outline:

- Composition and P/T state of interiors
- How materials deform at these conditions (timescales, mechanisms, T/T_m)
- Overview of lab studies
- Comparison of icy vs. rocky worlds (scaling)
- How material properties influence global-scale properties/dynamics





Interiors: Earth

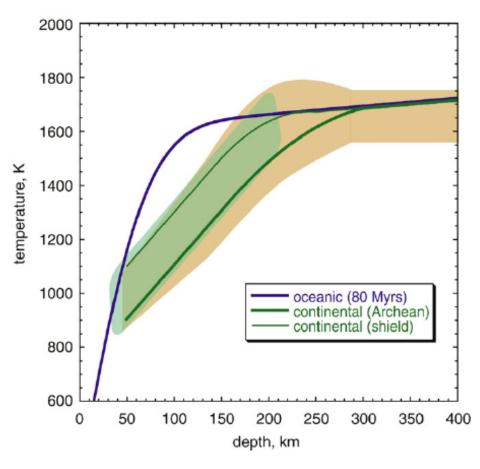
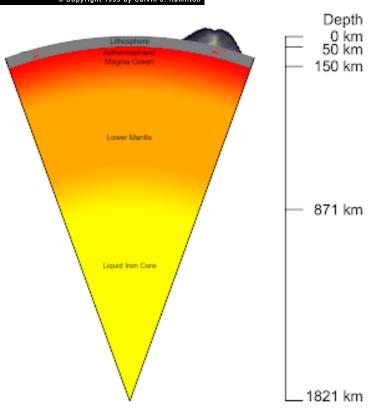


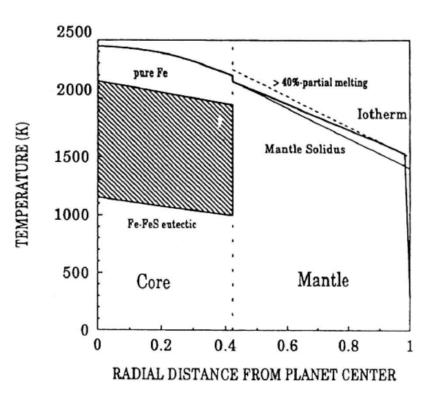
Fig. 2. The temperature-depth profiles for a typical old oceanic upper mantle (age of 80 Myrs) and for continents. For continents, two geotherms are shown corresponding to typical shield and Archean craton. Geotherms below 100 km are considered in the present study. The green region represents the temperature-depth data from mantle xenoliths, and the orange region corresponds to the temperature-depth profiles inferred from the inversion of seismic surface wave data from continental upper mantle.

[Karato, 2010]

The Interior of Io © Copyright 1999 by Calvin J. Hamilton

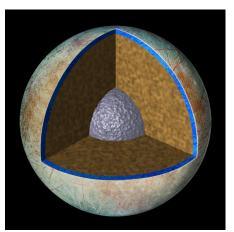
Interiors: lo



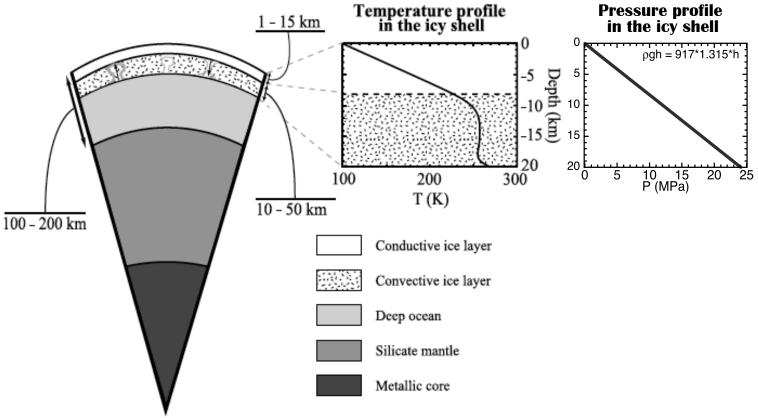


Thermal structure of Io from Schubert et al.

Mantle is dry peridotite very near to Tm



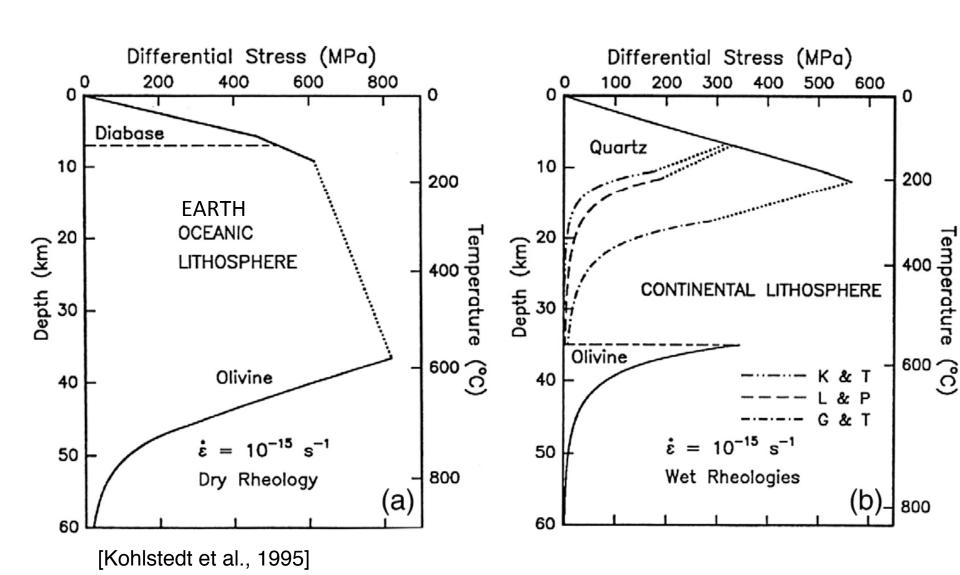
Interiors: Europa



[Tobie et al., 2003]

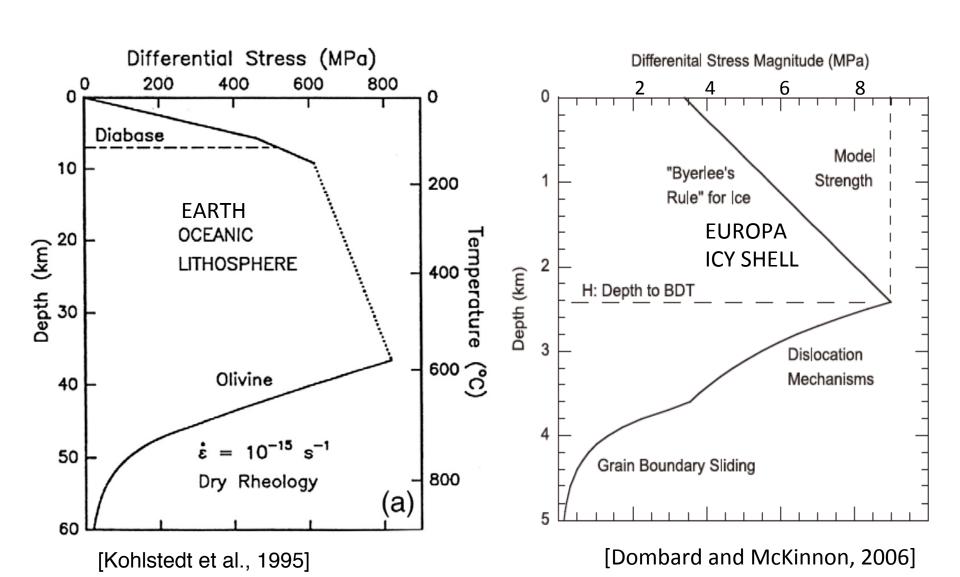
Strength envelopes:

From rheology of relevant minerals and T,P profiles



Strength envelopes:

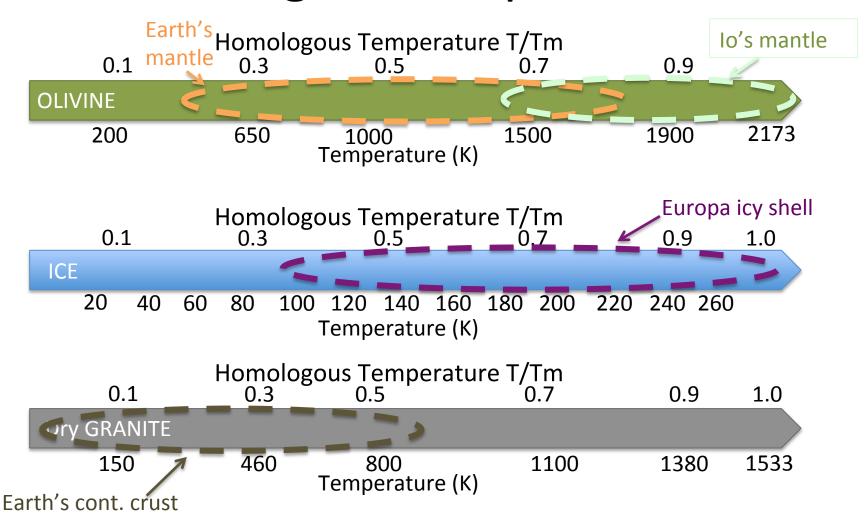
From rheology of relevant minerals and T,P profiles



Homologous temperature

0.1	Homologo 0.3	ous Temperatur 0.5	e T/Tm 0.7	0.9	
OLIVINE					
200	650	1000 Temperature (K)	1500	1900	2173
0.1	Homologo 0.3	ous Temperatur 0.5	e T/Tm 0.7	0.9	1.0
ICE					
20 40	60 80 100	120 140 160 Temperature (K)	180 200	220 240 260	
Homologous Temperature T/Tm					
0.1	0.3	0.5	0.7	0.9	1.0
Dry GRANITE					
150	460	800 Temperature (K)	1100	1380	1533

Homologous temperature



Deformation of polycrystalline materials occurs by motion of:

grain size

1D: point defects

2D: dislocations

3D: grain boundaries

(more than one happening at any given time. under different conditions and timescales one or more may dominate. They have different "signatures".)

ALSO influenced by: melt and second phases

A polycrystalline viscoelastic solid, warts and all

second

phases,

porosity

melt

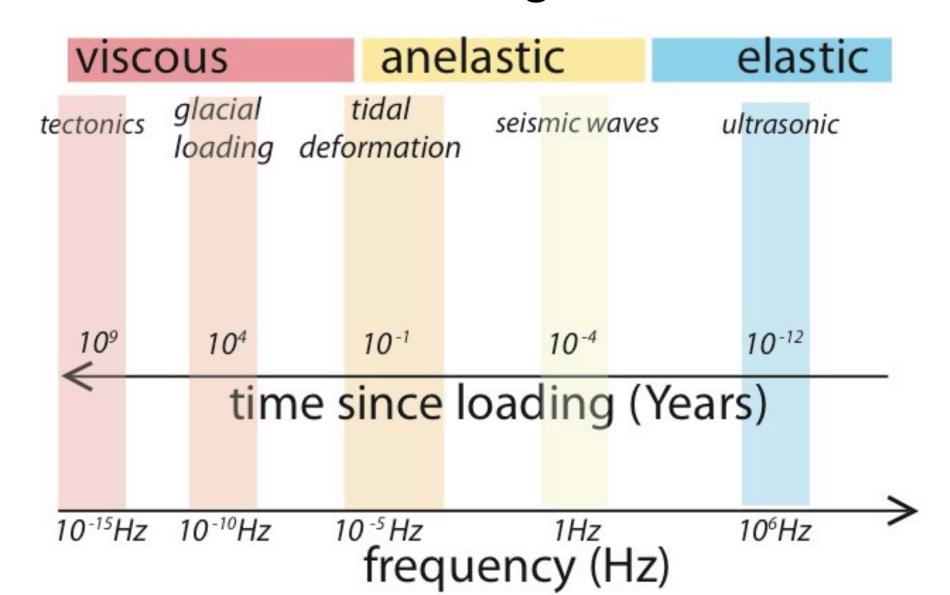
time,

freq.

Dislocations.

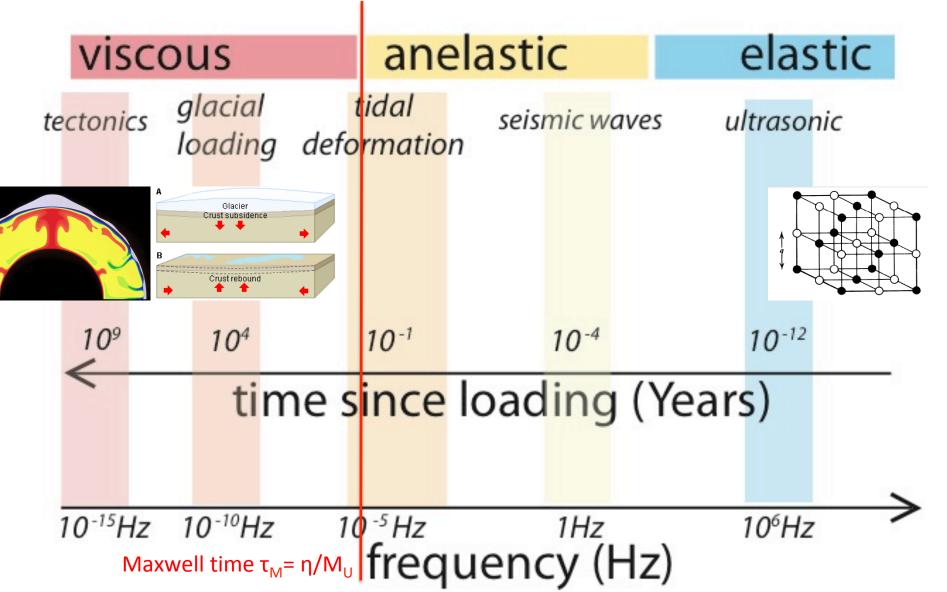
fabric

Viscoelasticity: Deformation at a range of time scales



Viscoelasticity:

Deformation at a range of time scales



Viscoelasticity

How do we measure viscosity and elasticity in the lab?

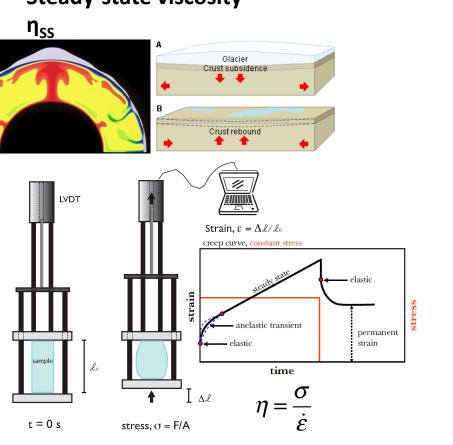
viscous

anelastic

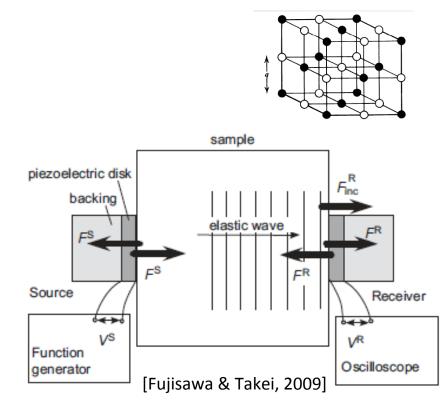
elastic

Creep experiments

Steady-state viscosity



ultrasonics **Elastic Modulus** *k* **or E**_u



Viscoelasticity

Viscous behavior; strain rate is proportional to stress:

$$\sigma = \eta \dot{\varepsilon}$$

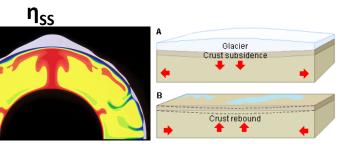
Steady-state viscosity



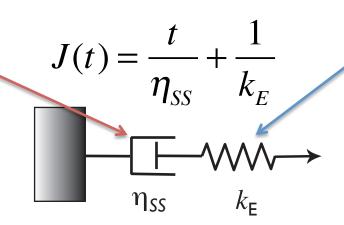
Elastic behavior is instantaneous elasticity and instantaneous recovery. Follows Hooke's Law:

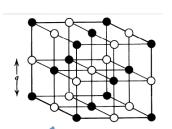
$$\sigma = E \varepsilon$$

Elastic Modulus k or E

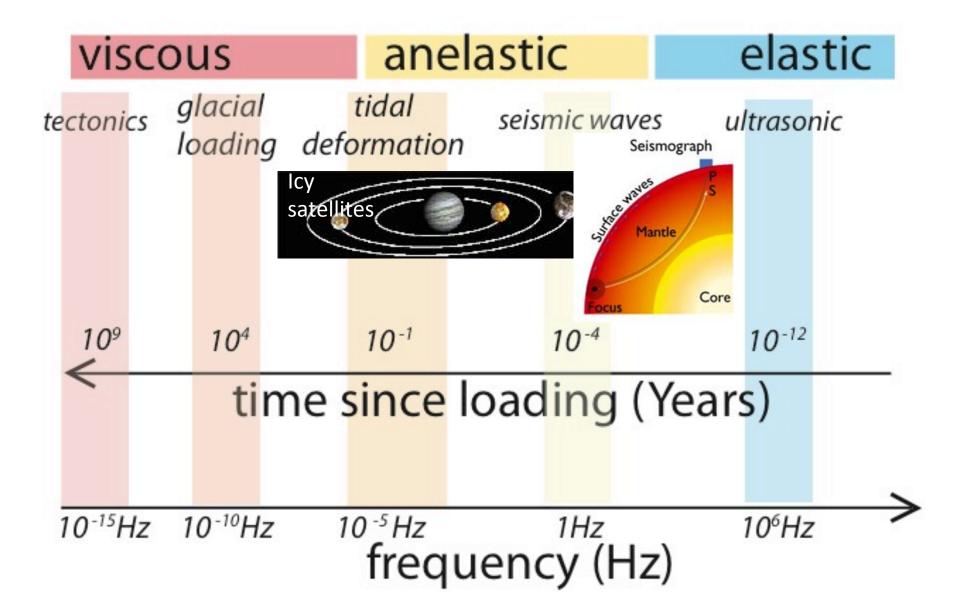


Simplest form of viscoelasticity is the Maxwell model:

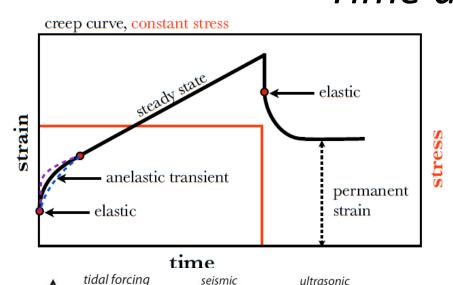




Viscoelasticity: in between the two extremes?



Viscoelasticity Time domain



$$J(t) = \frac{\varepsilon(t)}{\sigma} = \frac{t}{\eta_{SS}} + F(t) + \frac{1}{k_E}$$

Creep compliance



Laplace transform



Complex compliance

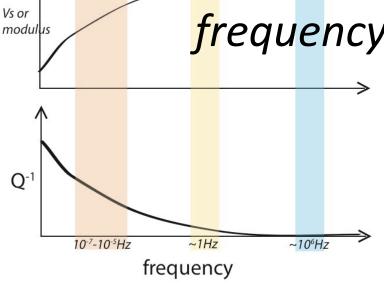
$$\sigma(t) = \sigma_0 \exp(i\omega t); \varepsilon(t) = \sigma_0 J^* \exp(i\omega t)$$

$$J * (\omega) = J_1(\omega) + iJ_2(\omega)$$

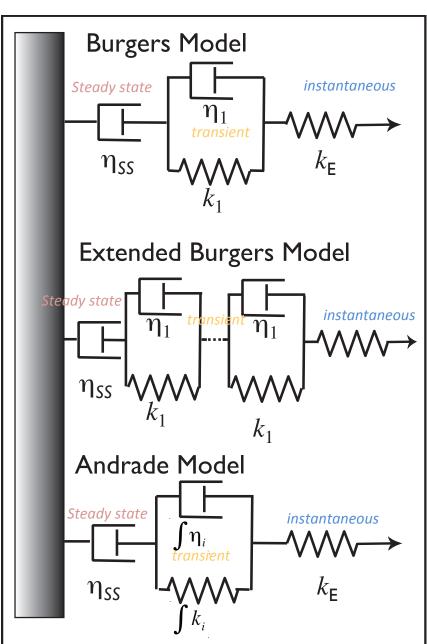
$$\omega = 2\pi f$$

$$Q^{-1} = \frac{J_2}{J_1} \quad \text{-Loss compliance} \\ \text{-Storage compliance}$$

$$E = \frac{1}{\sqrt{J_1^2 + J_2^2}}$$



Mechanical models for dissipation

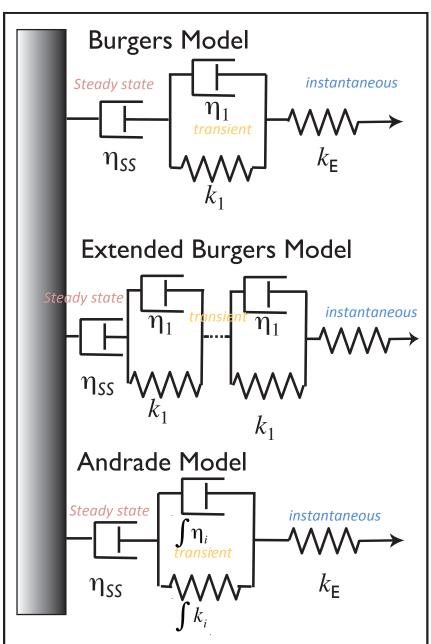


$$J(t)^{Burgers} = \frac{\varepsilon(t)}{\sigma} = \frac{t}{\eta_{SS}} + \frac{1}{k_2} \left[1 - \exp\left(\frac{-t}{\tau}\right) \right] + \frac{1}{k_E}$$

$$J(t) = \frac{\varepsilon(t)}{\sigma} = \frac{t}{\eta_{SS}} + \frac{1}{k_E} \Delta \int_{\tau_m}^{\tau_M} \left[1 - e^{-t/\tau} \right] D(\tau) d\tau + \frac{1}{k_E}$$

$$J(t)^{And} = \frac{\mathcal{E}(t)}{\sigma} = \frac{t}{\eta_{SS}} + \beta t^m + \frac{1}{k_E}$$

Mechanical models for dissipation



$$J_1^B(\omega) = \frac{1}{k_E} + \frac{k_2}{k_2^2 + \eta_2^2 \omega^2}$$
$$J_2^B(\omega) = \frac{\eta_2 \omega}{k_2^2 + \eta_2^2 \omega^2} + \frac{1}{\eta_{ss} \omega}$$

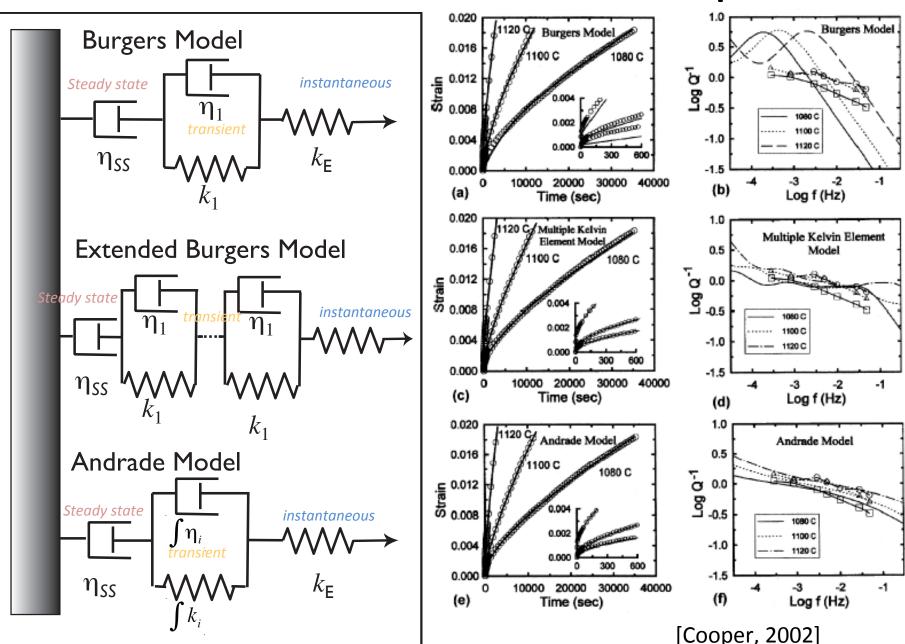
$$J_1^{EB}(\omega) = \frac{1}{k_E} \left[1 + \Delta \int_0^{\infty} D(\tau) d\tau / (1 + \omega^2 \tau^2) \right]$$

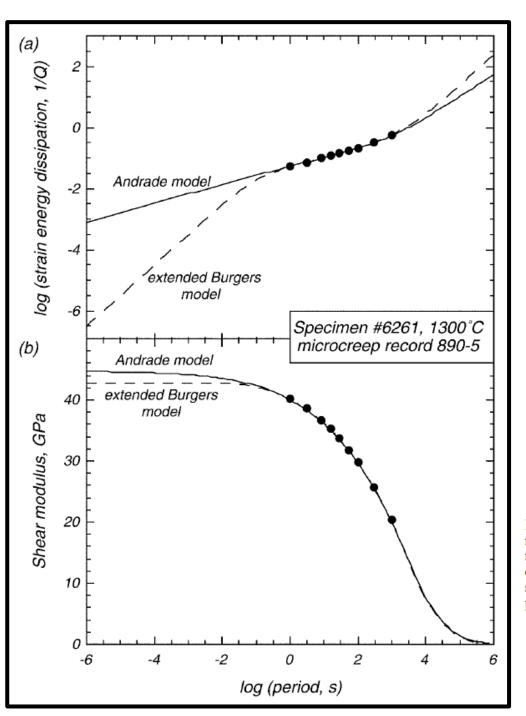
$$J_2^{EB}(\omega) = \frac{\omega}{k_E} \Delta \int_0^{\infty} \tau D(\tau) / (1 + \omega^2 \tau^2) + \frac{1}{\eta_{SS} \omega}$$

$$D(\tau) = \frac{\alpha \tau^{-(1-\alpha)} H(\tau - \tau_m) H(\tau_M - \tau)}{\tau_M^{\alpha} - \tau_m^{\alpha}}$$

$$J_1^{And}(\omega) = \frac{1}{k_E} + \beta \Gamma(1+m)\omega^{-m} \cos\left(\frac{m\pi}{2}\right)$$
$$J_2^{And}(\omega) = \beta \Gamma(1+m)\omega^{-m} \sin\left(\frac{m\pi}{2}\right) + \frac{1}{\eta_{SS}\omega}$$

Mechanical models for dissipation





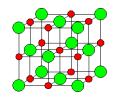
Burgers vs. Andrade: work equally well

Fig. 7a, b Andrade and extended Burgers model fits (Table 2) to a the modulus dispersion and b the dissipation information extracted from microcreep record 890-5 at 1300 °C. The alternative rheologies fit the data (indicated by the *plotting symbol*) equally well but diverge substantially in the description of Q^{-1} at much shorter and longer periods

[Tan, Jackson and Fitz Gerald, 2001]

Deformation of polycrystalline materials occurs by motion of:

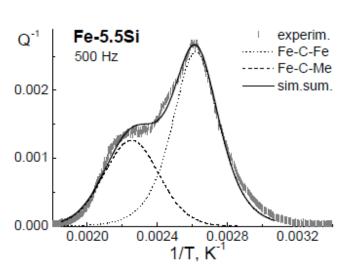
• 1D: point defects



• 2D: dislocations

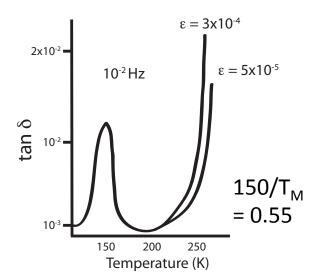
• 3D: grain boundaries

Distinct length and timescales gives a peak in spectra.
Point defect motion makes peaks.



Snoek-type and Zener relaxation in Fe - Si - Al alloys

I. S. Golovin^{1,a}, S. Jäger², V. A. Semin¹, G. V. Serzhantova¹, H.-R. Sinning², O. A. Sokolova¹, F. Stein³, S. A. Golovin¹



Attenuation in single crystal ice with Debye peak at ~150K due to proton rearrangement [after Tatibouet et al., 1981]

Deformation of polycrystalline materials occurs by motion of:

Distinct length and timescales

Grain boundary diffusion displays

gives a peak in spectra.

a broad distribution.

• 1D: point defects

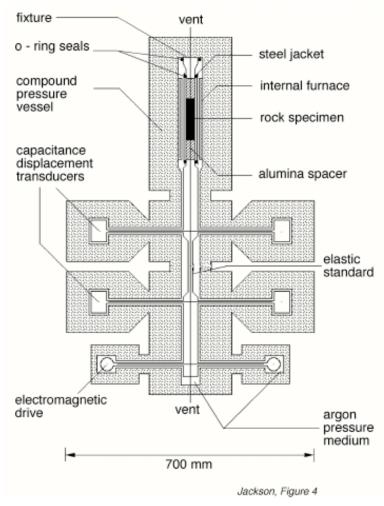
2D: dislocations

3D: grain boundaries

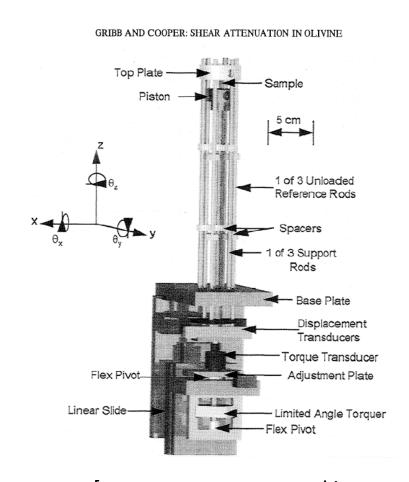
Local stress Local slip at concentration Global slip (creep) No slip planer segments and relaxation Grain **boundary**

How can we identify the mechanism of attenuation in experiments?

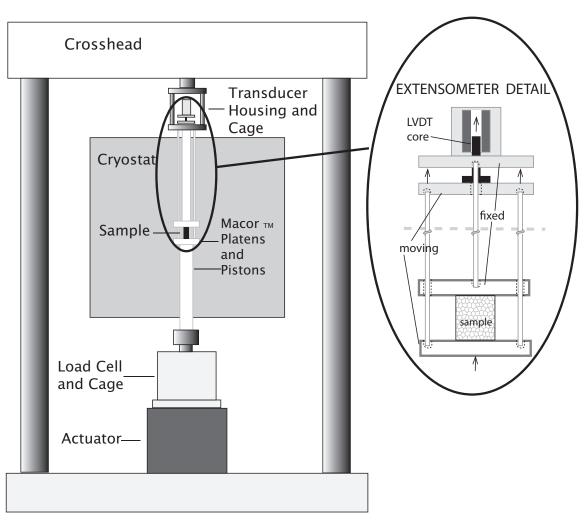
- Look at the spectra. Are there peaks (i.e. narrow relaxation times) or a broad bands (distribution of relax. time)?
- Look at the apparent thermal and grain size dependences. Can you find internal scaling? How does the activation energy and gs-dependence compare to steady-state processes?
- Look at the microstructure.



[custom apparatus used in many Jackson, Faul, Farla papers; described in Jackson and Paterson, 1993]



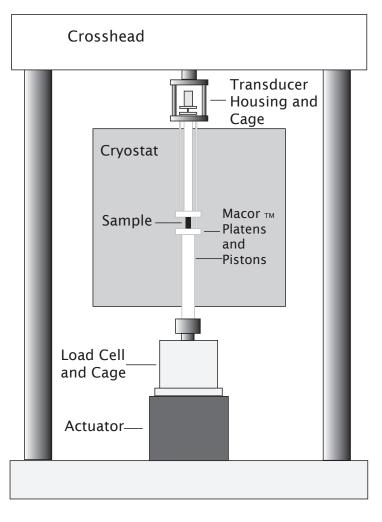
[custom apparatus used in Gribb and Cooper, 1998; 2000; Sundberg and Cooper, 2010]



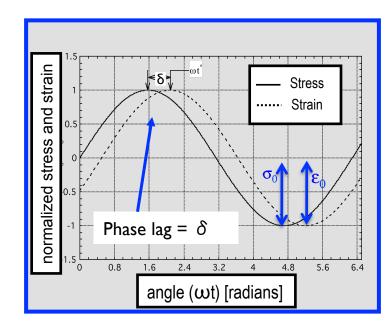


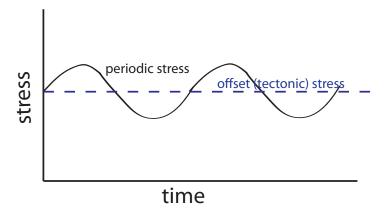
Large standing Instron with cryostat used at Brown

Table top Instron with cryostat used at JPL

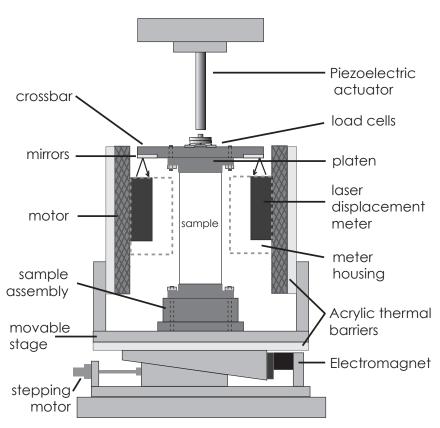


$$Q^{-1} = \tan \delta = \frac{J_2}{J_1};$$
$$E = \left(\frac{\sigma_0}{\varepsilon_0}\right)$$

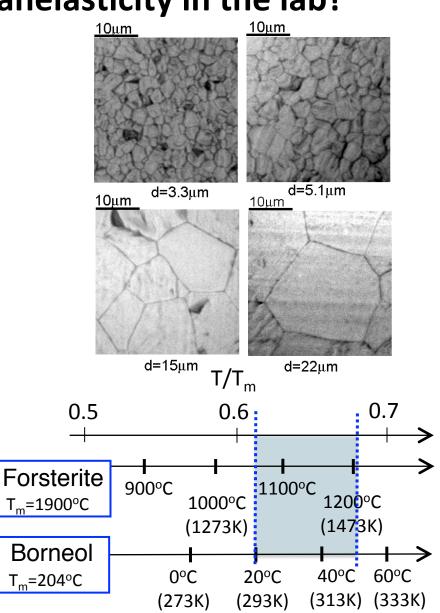




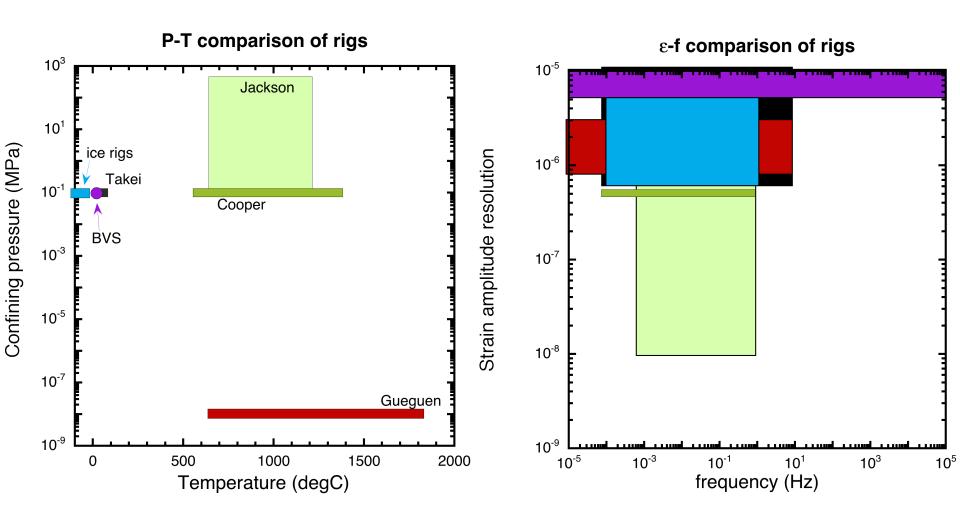
Control grain size, T, impurities etc.



Analogue samples: borneol (C₁₀H₁₈O) [Takei, Fujisawa, McCarthy, JGR116, 2011]

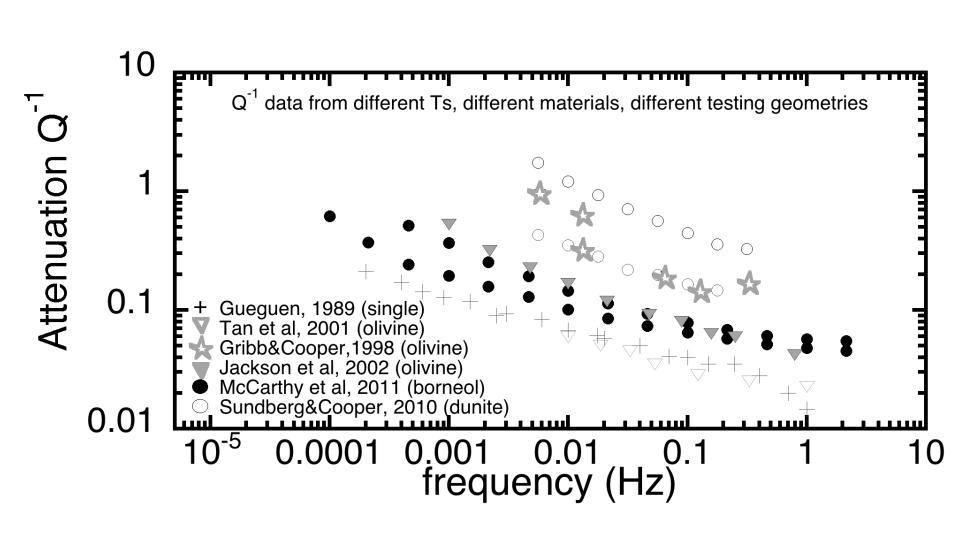


Comparison of forced oscillation rigs

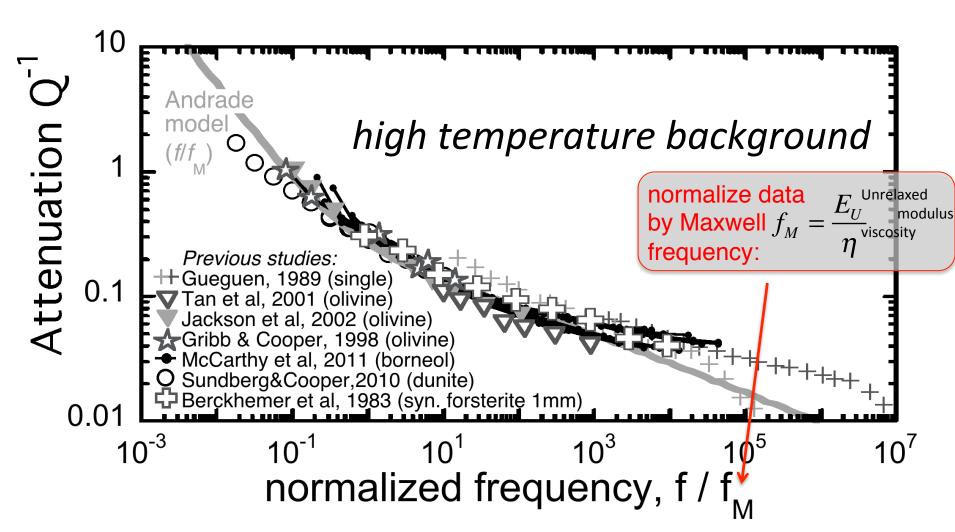


Empirical Observations:

comparison of multiple studies



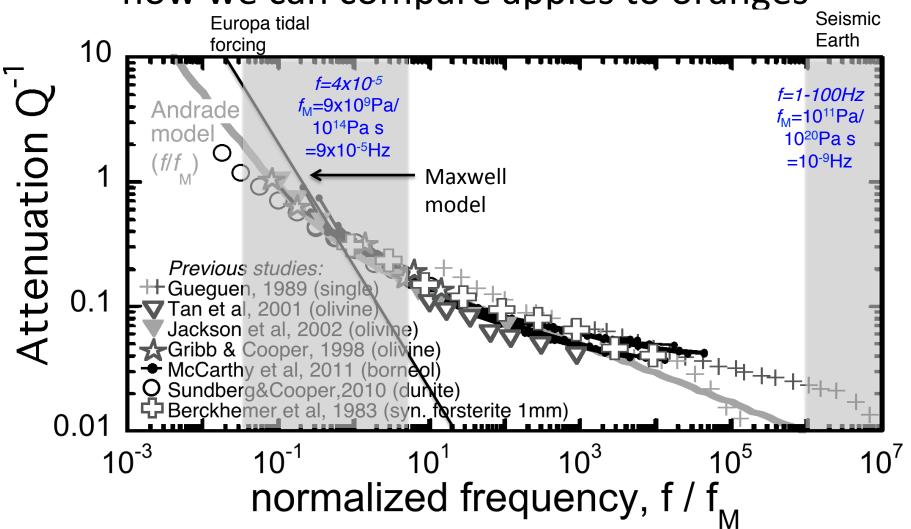
Empirical Observations: Normalization by Maxwell frequency comparison of multiple studies



Empirical Observations:

Normalization by Maxwell frequency

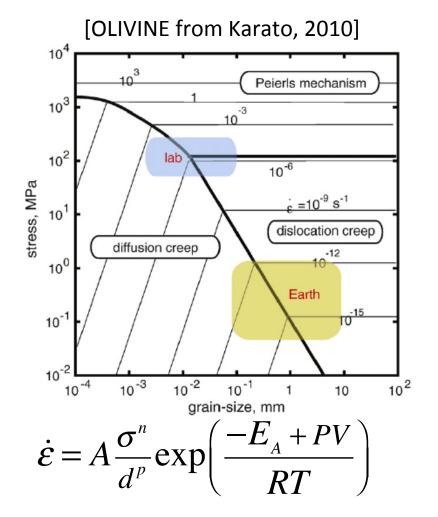
how we can compare apples to oranges



Maxwell curve approach: take a description of the HTB and normalize by Maxwell frequency for the material or planetary setting of your choice

$$\frac{1}{f_{M}} = \tau_{M}(d, T, P) = J_{U}(T, P) \cdot \eta_{0} \left(\frac{d}{d_{r}}\right)^{-p} \exp\left[\frac{U}{R}\left(\frac{1}{T} - \frac{1}{T_{r}}\right)\right] \exp\left[\frac{V}{R}\left(\frac{P}{T} - \frac{P_{r}}{T_{r}}\right)\right]$$

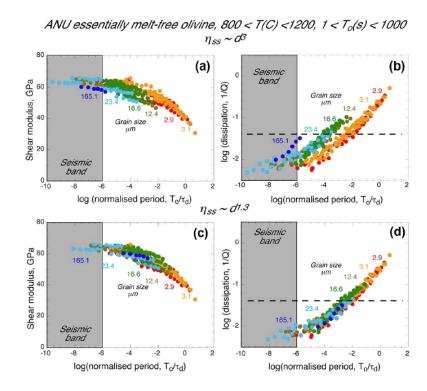
Moduli and viscosities of Earth and planetary materials pretty well known



[ICE from Prieto-Ballesteros et al., 2010] 10^{2} T=220 K Durham et al. Dislocation creep n=4Goldsby and 10^{0} stress, MPa Kohlstedt Icy satellites 10-2 Superplastic Icy satellites Glaciers creep n=1.8Ice sheets 10-4 .001.010.110 100 grain size, mm

The fine print: problems with Master curve approach

- Some disagreement about d-dependence
- A peak at high frequency overlays the HTB
- Melt effect on GBS known, but "squirt" not well known
- Water effect not well known
- Dislocation effect not well const



$$\tau_{M}(d,T,P) = J_{U}(T,P) \cdot \eta_{0} \left(\frac{d}{d_{r}}\right)^{-p} \exp\left[\frac{U}{R}\left(\frac{1}{T} - \frac{1}{T_{r}}\right)\right] \exp\left[\frac{V}{R}\left(\frac{P}{T} - \frac{P_{r}}{T_{r}}\right)\right]$$
[Jackson, Faul, Skelton, 2013]

The fine print: problems with Master curve approach

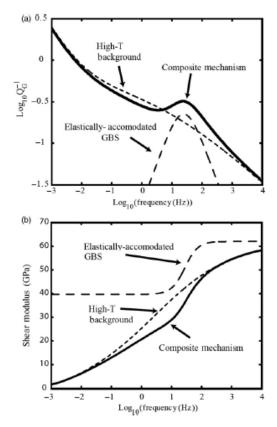
- Some disagreement about d-dependence
- A peak at high frequency overlays the HTB

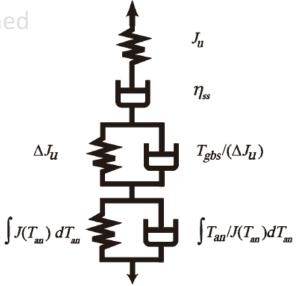
Melt effect on GBS known, but "squirt" not well known

Composite creep function

Water effect not well known

Dislocation effect not well constrained





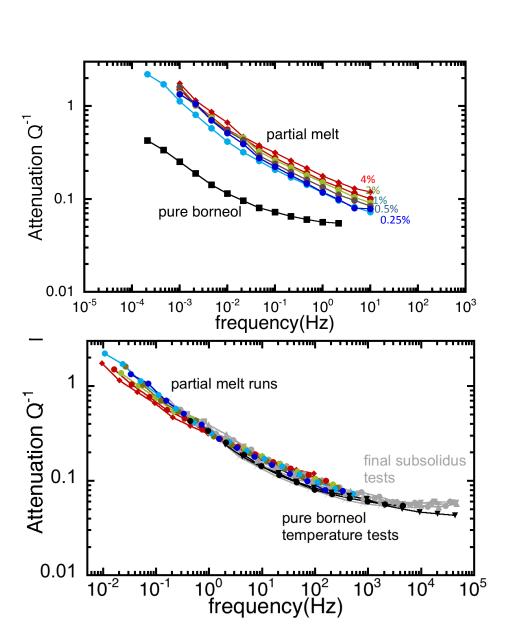
- A. Unrelaxed compliance
- B. Steady-state creep
- C. Elastically-accomodated GBS; exponential decay with discrete relaxation time
- D. High-temperature background; power law decay with continuous distribution of relaxation times

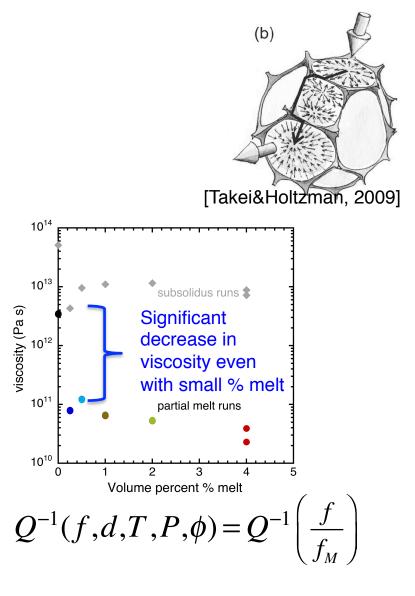
$$J(t) = J_U + \frac{t}{\eta_{ss}} + \left((J_U \Delta)(1 - e^{\frac{-t}{T_{gbs}}}) \right) + At^n$$
$$= A + B + C + D$$

M. Sundberg and R.F. Cooper

[Sundberg and Cooper, 2010]

Effect of melt on GBS and master curve





[McCarthy and Takei, GRL38, 2011]

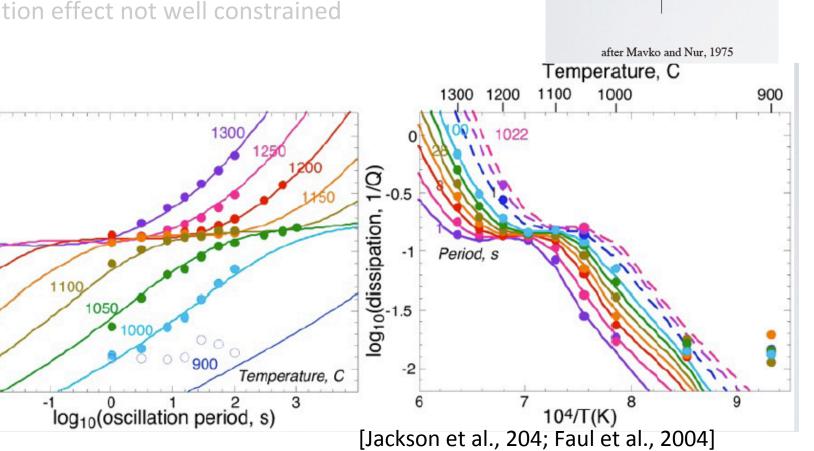
The fine print: problems with Master

curve approach

- Some disagreement about d-dependence
- A peak at high frequency overlays the HTB
- Melt effect on GBS known, but "squirt" not well known
- Water effect not well known

log₁₀(dissipation, 1/Q)

Dislocation effect not well constrained



Melt squirt:

pressure driven melt flow

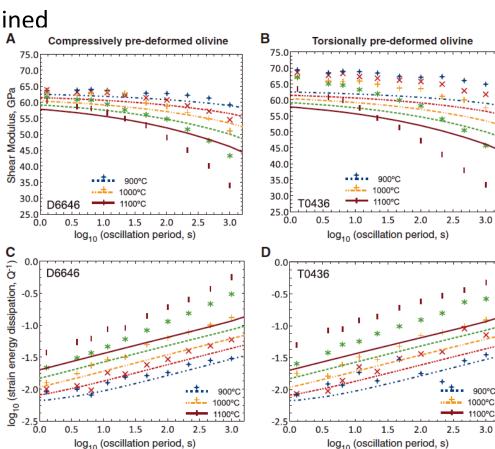
The fine print: problems with Master curve approach

- Some disagreement about d-dependence
- A peak at high frequency overlays the HTB
- Melt effect on GBS known, but "squirt" not well known
- Water effect not well known
- Dislocation effect not well constrained

This study used samples pre-deformed in the dislocation creep regime (longitudinal and torsional)

In both cases, the Q⁻¹ was greater than the prediction (of diffusion-GBS)

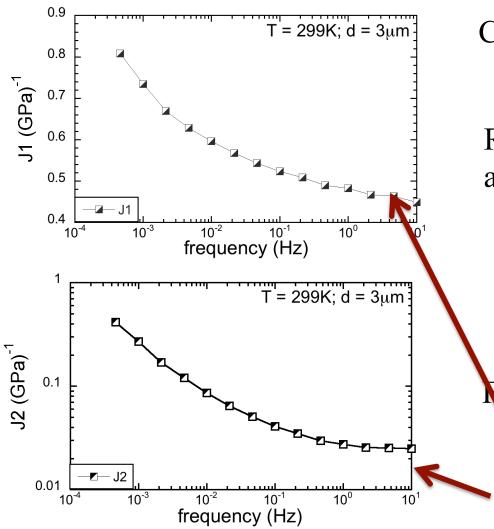
[Farla et al., 2012]



Defining J1 and J2

$$\sigma(t) = \sigma_0 e^{i\omega t}$$

$$\varepsilon(t) = \sigma_0 J^* e^{i\omega t}$$



Complex compliance

$$J^*(\omega) = J_1 - iJ_2$$

Relation between Q and E

and
$$J_1$$
 and J_2

$$Q^{-1} = \frac{J_2}{J_1}$$

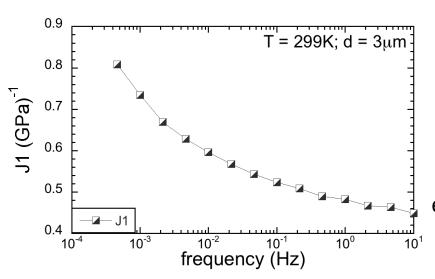
$$E = \frac{1}{\sqrt{J_1^2 + J_2^2}}$$

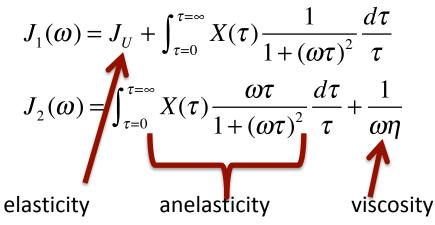
From Q and E to J1 and J2

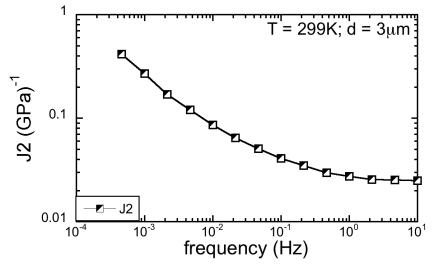
$$J_1(\omega) = \frac{1}{E(\omega)\sqrt{1 + (Q^{-1}(\omega))^2}}$$

$$J_2(\omega) = Q^{-1}(\omega)J_1(\omega)$$

Defining J1 and J2





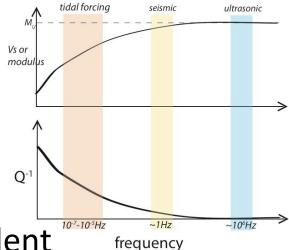


Well approximated by:

$$\begin{split} J_{1}(P) + \frac{P}{\pi^{2}\eta} &= J_{1}(P_{R}) + \frac{2}{\pi} \int_{\ln P_{R}}^{\ln P} J_{2} d\ln P \\ J_{2}(\omega) &= \frac{\pi}{2} X \left(\tau - \frac{P}{2\pi} \right) + \frac{P}{2\pi\eta} \end{split}$$

From Nowick and Berry

Takeaways



- Material properties are frequency dependent
- Anelastic behavior depends on mechanism (defects)
- Mechanical models can describe behavior of materials (Andrade, Burgers etc.)
- Although "apparent" relationships of Q⁻¹ on d, T, P, Φ etc. exist, these are related through viscosity, so use Maxwell freq. scaling to compare apples to oranges.
- The "high temperature background" observed in labs is also observed in seismic studies (absorption band). Similitude in experiments suggests same mechanism (GBS). (some of the nuanced differences may come in the "fine print")